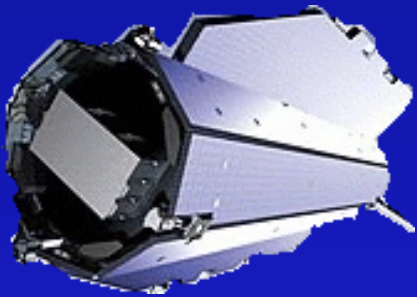


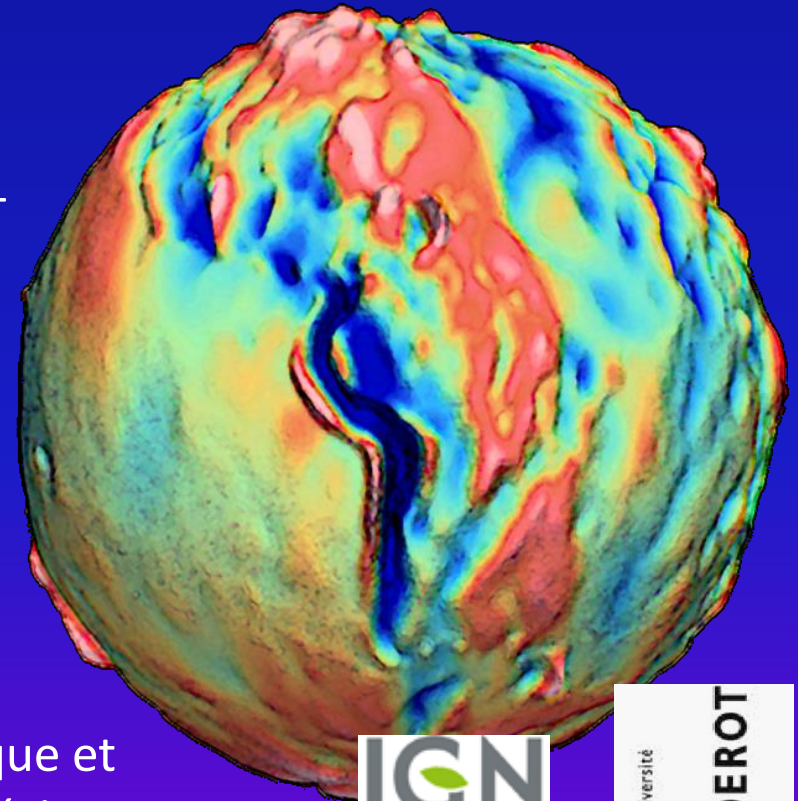
# Satellite gravity: a probe on Earth's system dynamics



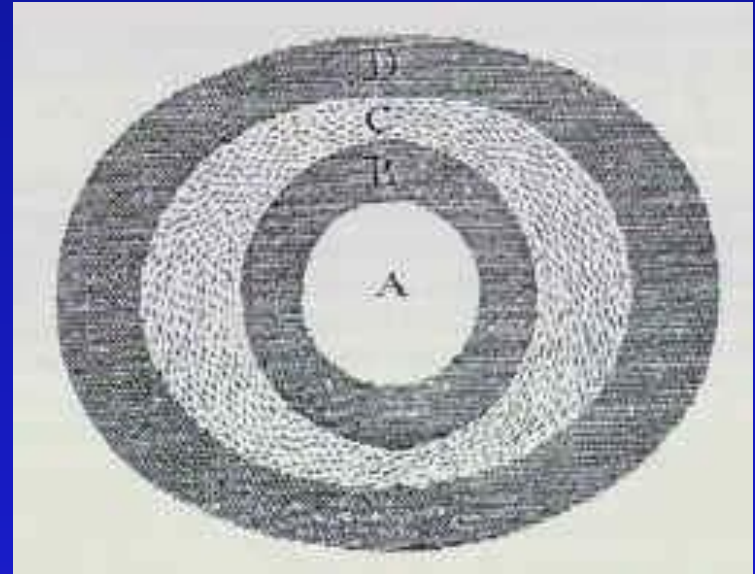
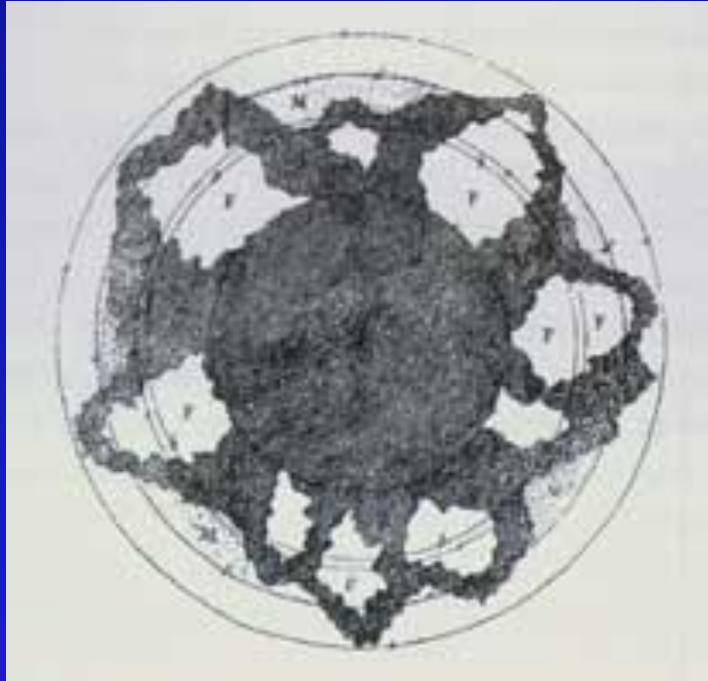
Isabelle Panet<sup>1</sup>

Merci à G. Pajot-Métivier, M. Greff,  
L. Métivier, M. Diament, M. Mandeau

<sup>1</sup> Institut National de l'Information Géographique et  
Forestière, Laboratoire de Recherche en Géodésie,  
Université Paris Diderot



- Early measurements of our planet's gravity are related to questions about its shape and interior

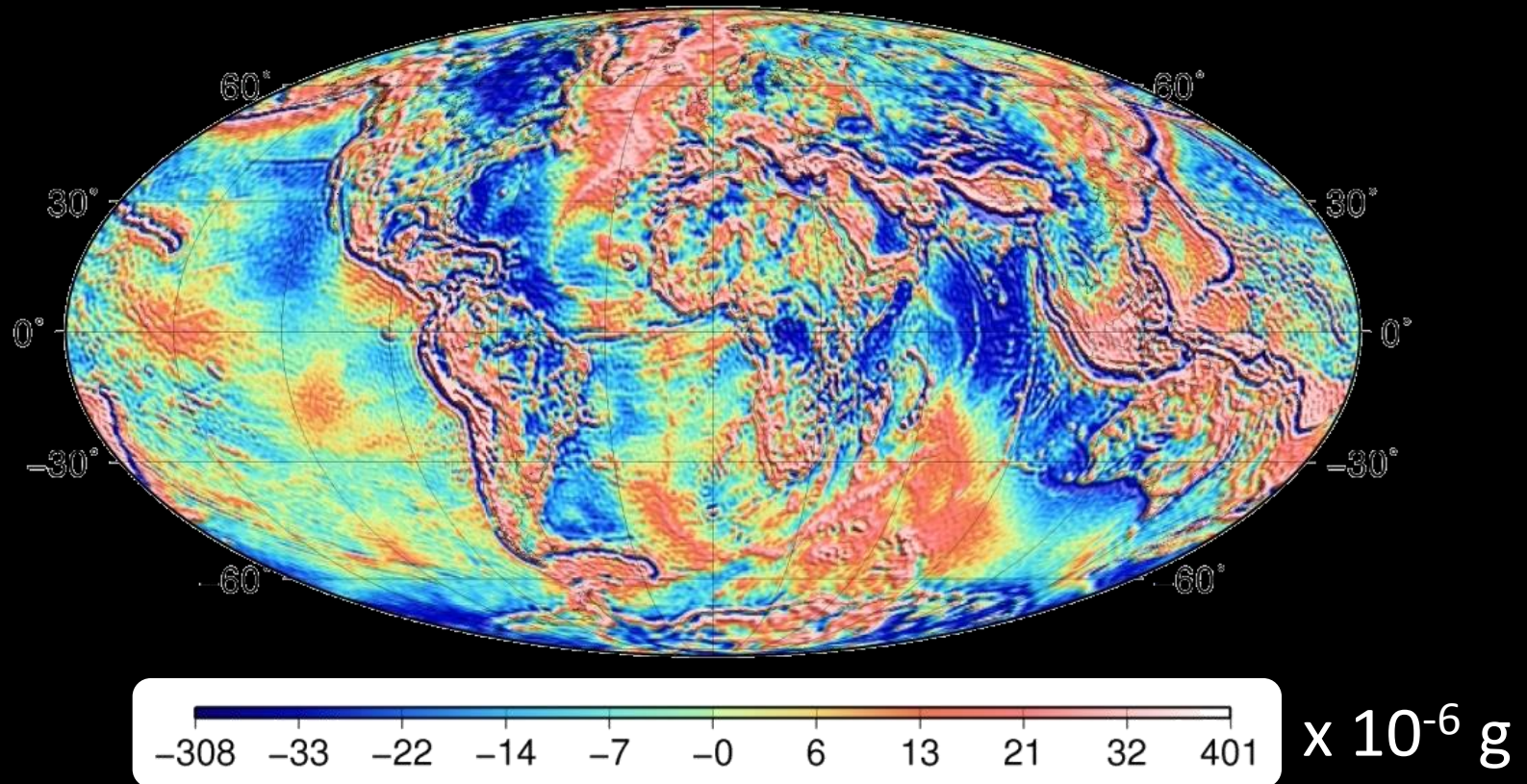


???

*Source: web ENS Lyon*

- Cavendish (18th century): the torsion balance as a precursor to modern measurements of gravity gradients

# Earth's gravity intensity varies in space...

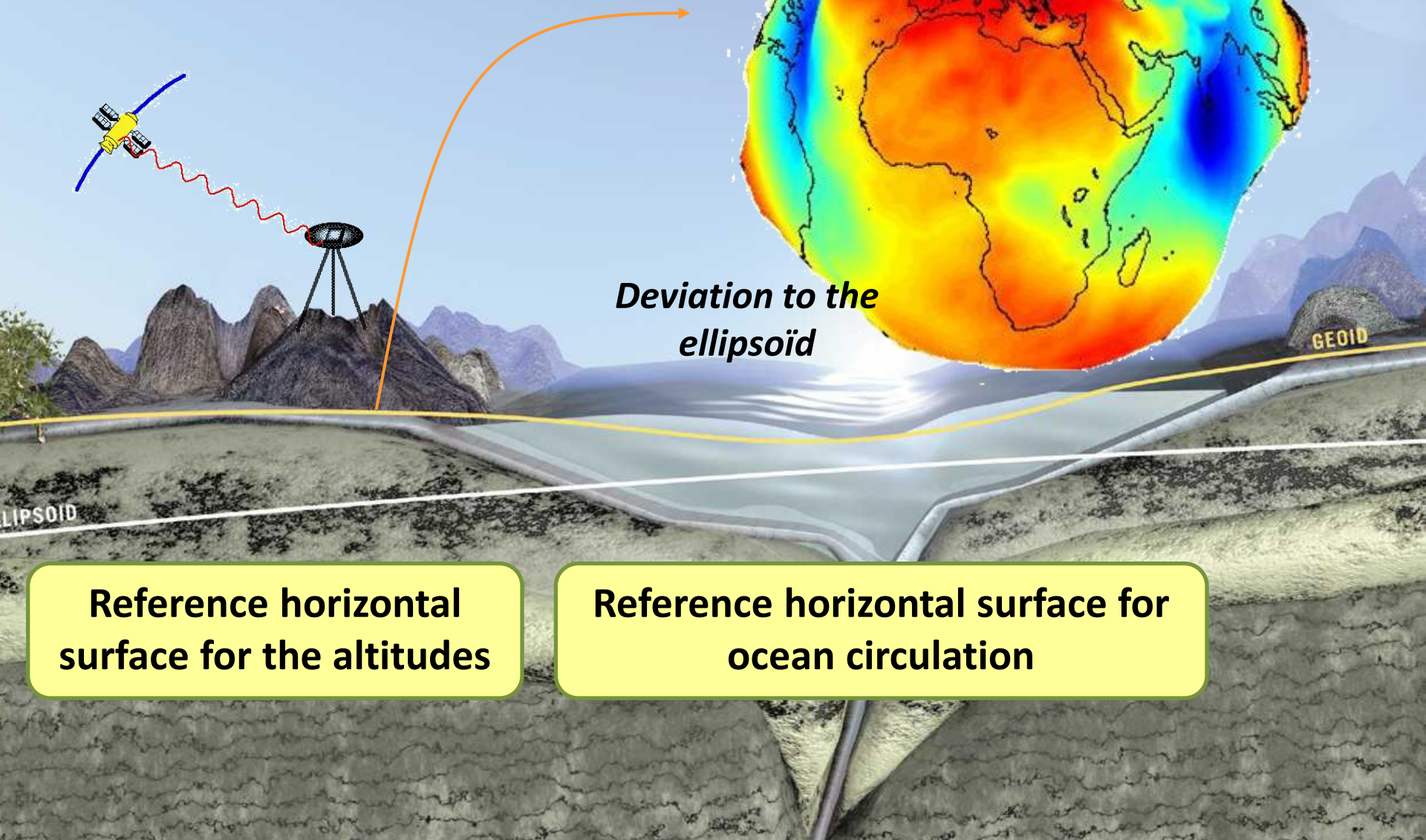


*contribution from a homogeneous ellipsoidal Earth removed*

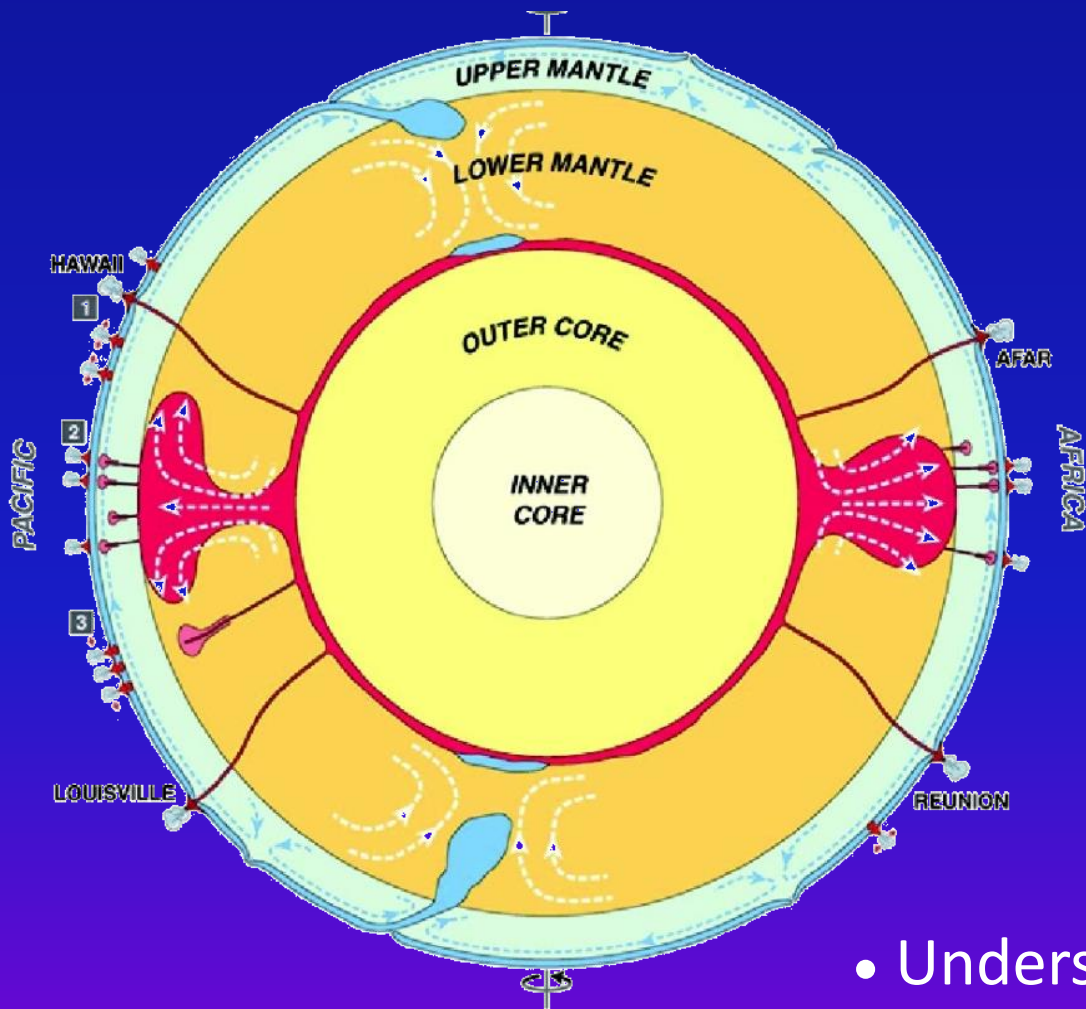


# Earth's geoid gives the horizontal

$\pm 100$  m



# Gravity varies as Earth's mass distribution is not homogeneous



*Courtilot et al. (2003)*

- The rigid crust floats on a mantle which behaves as a highly viscous fluid at « long » time scales.
- The internal convection releases Earth's internal heat to the surface and the outer space
- Understand current structure and Earth's evolution in time??



# Gravity varies in time as these masses move

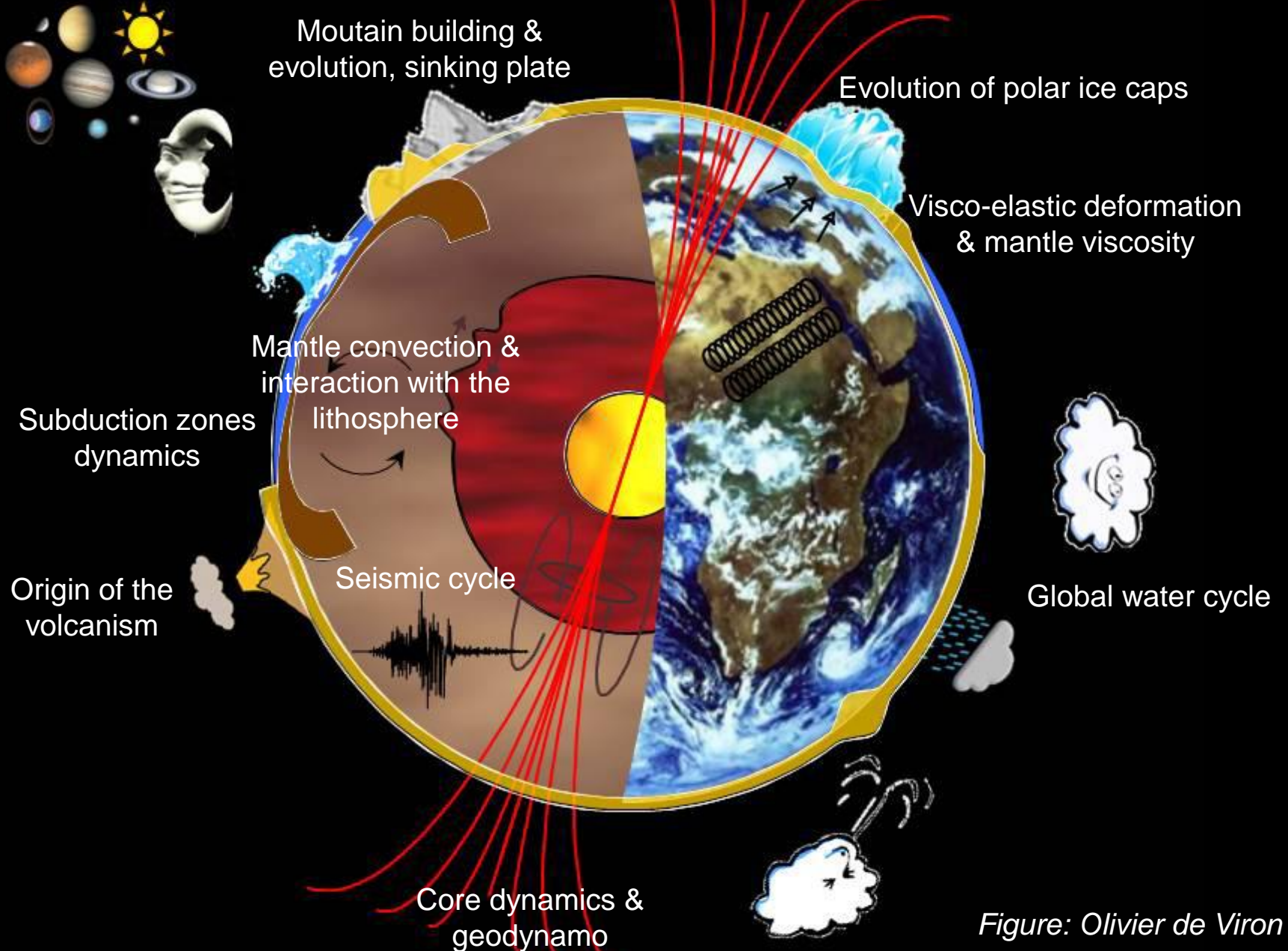


Figure: Olivier de Viron

- Even if differential measurements of gravity are an early concept (Cavendish, Eötvös), analyzing the field intensity is more usual

**Easier to measure**

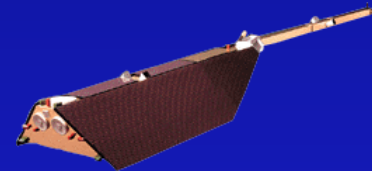
**Easier to interpret**

- However, separating signals from superimposed sources in gravity data is a crucial step, that benefits from a directional information

**Identify sources geometries**

*Great progress in knowing Earth's gravity comes from satellite gravity missions*

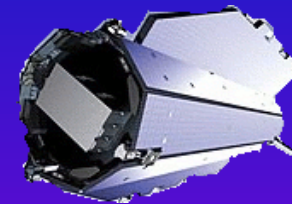
*CHAMP*  $g$



*GRACE*  $g(t)$

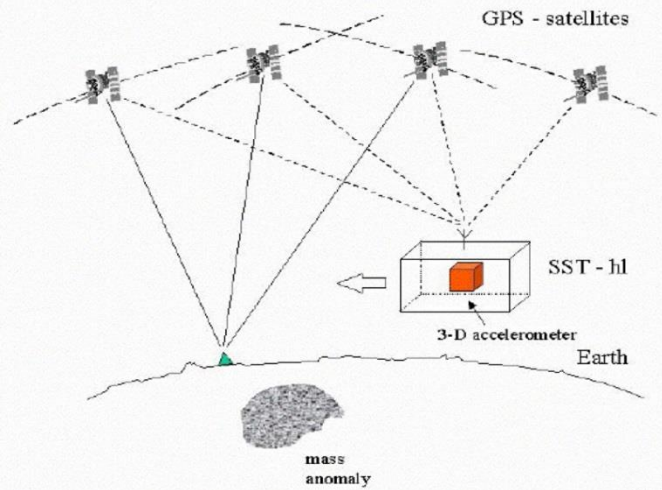


*GOCE*  $\vec{g}$

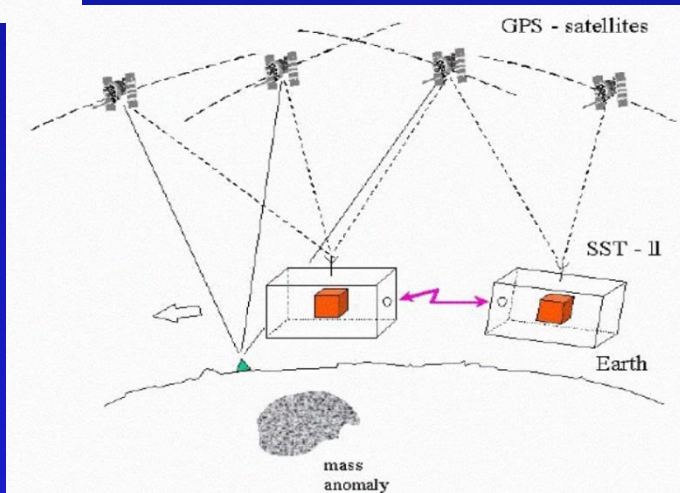




# Satellite gravity missions

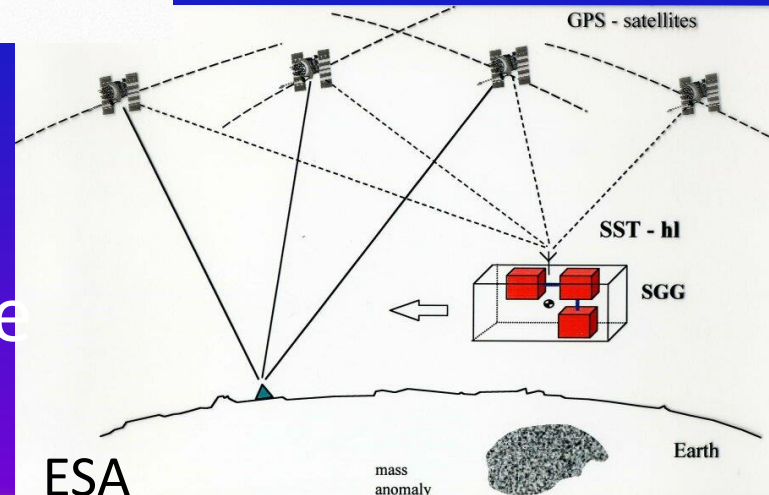


**CHAMP (2000-2010)**



**GRACE (2002- ...)**  
*Time variations*

**GOCE (2009-2013)**



ESA

- Lower and lower orbits  
GOCE: ~250 a 225 km altitude !
- Differentiating more and more  
Amplify details

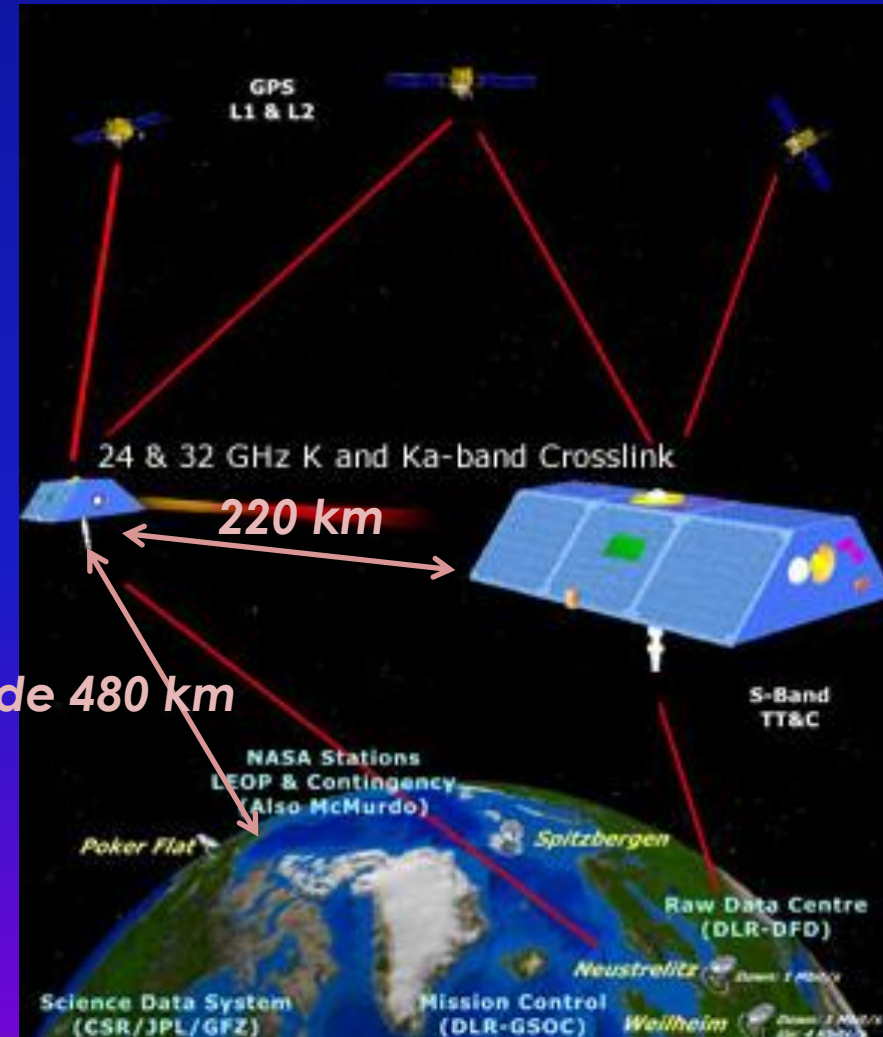
# GRACE

## *Gravity Recovery And Climate Experiment*

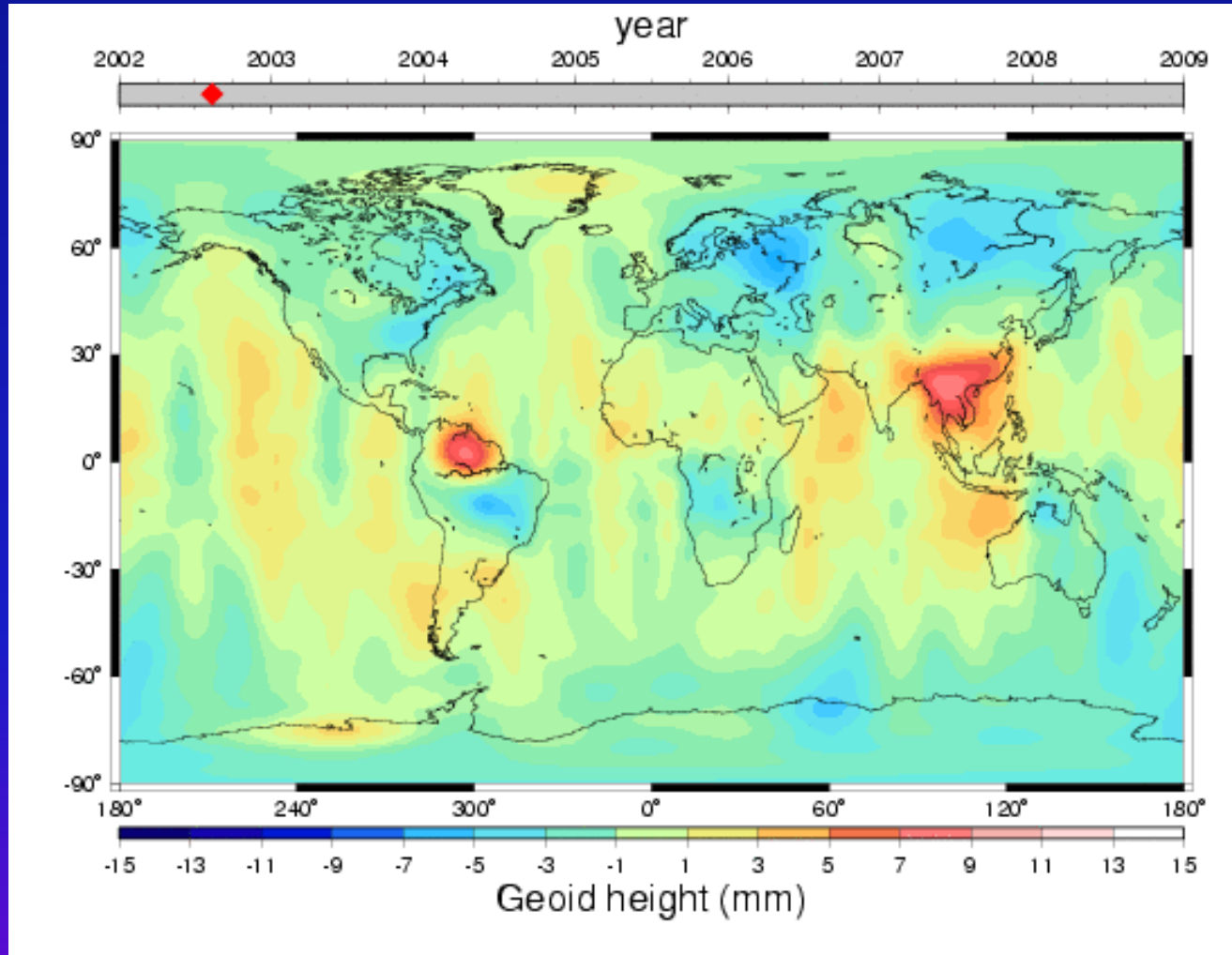
*Inter-satellite distance variation*

*Precision :  
microns / microns per second*

*« One arm gradiometer »*



# Temporal variations of Earth's geoid



$\pm 1.5$  cm

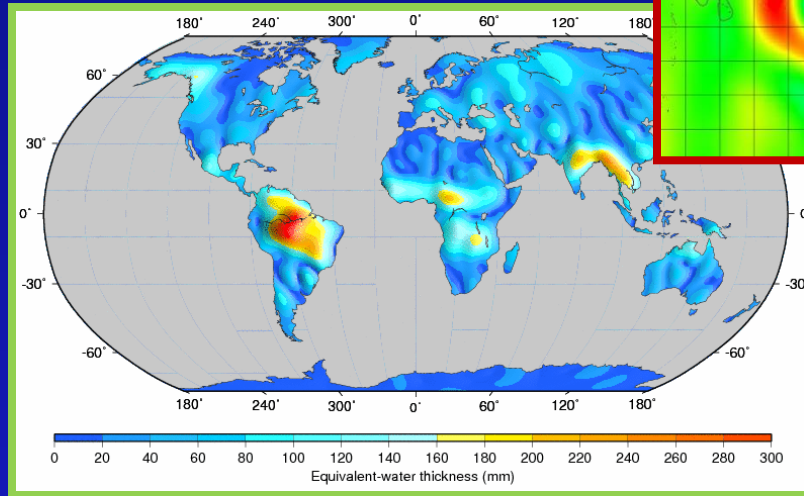
Biancale et al., CNES/GRGS



« instantaneous »

Time scale

a few mm

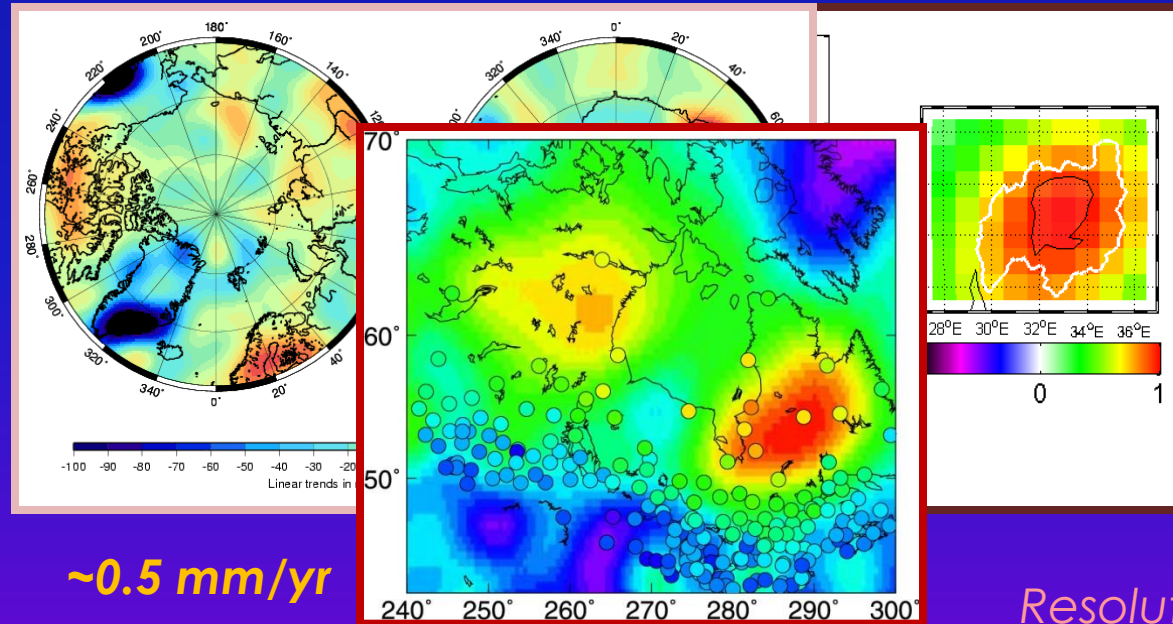


At GRACE scales...

seasonal

inter-annual

a few mm/yr



~0.5 mm/yr

Resolution (km)

et al.

4000

1200

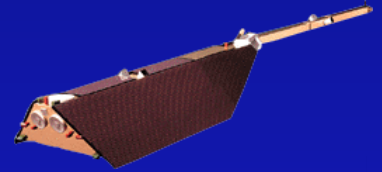
800

400

Ramillien, de Viron, Tregoning, Panet

# Satellite gravity missions

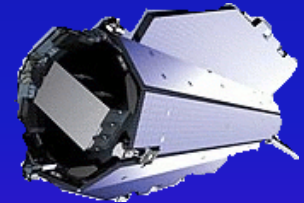
CHAMP  $g$



GRACE  $g(t)$



**GOCE**  $\vec{g}$



Rather than the geoid or the gravity intensity, let's look at the tiny variations of the gravity **vector**

# GOCE

- A very low altitude (255 km, lowered to ~225 km)

→ Limit field attenuation

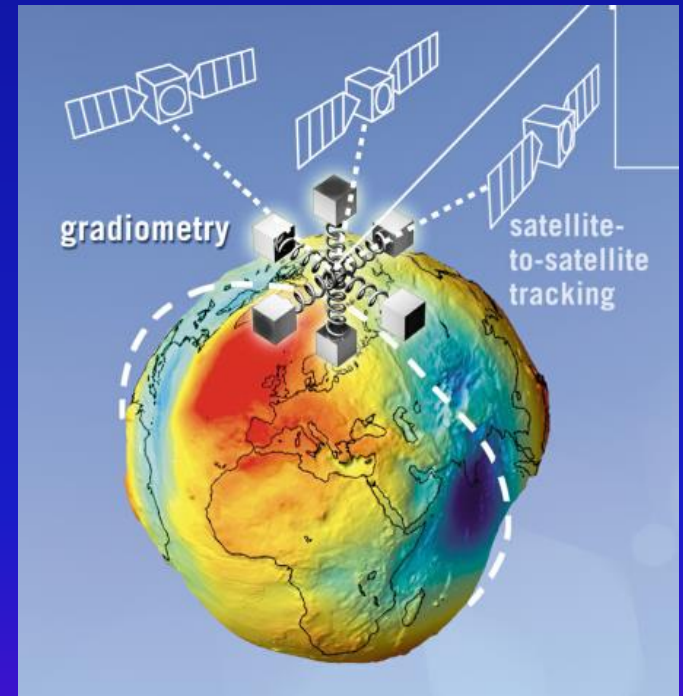
- Sensitivity to small structures

→ Amplification by differentiation:  
gradiometry.

→ Direction of measurement needs  
to be known as accurately

- Compensation of non-gravitational  
forces along the orbit

- Orbit determination: GPS + laser ranging





# At the core of the mission: mapping the geoid and the gravity field at high resolution from the gravity gradients

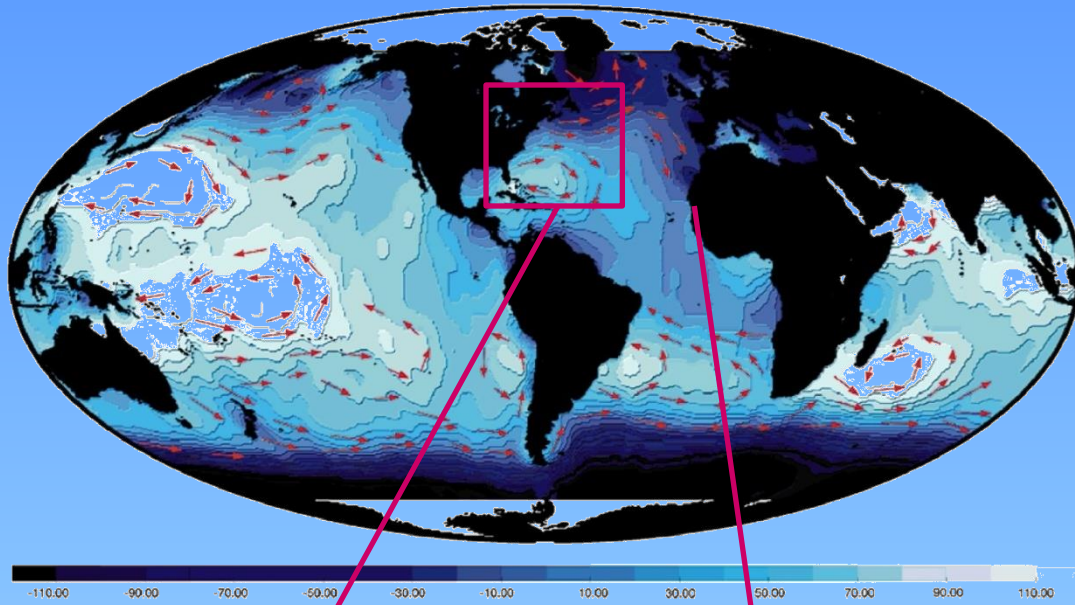
Objective: geoid with centimeter accuracy at 100 km resolution  
gravity anomalies with a  $10^{-6}$  g precision

*Ocean circulation*

*Altitudes détermination*

*Geometry of the bedrock below polar ice caps*

*Earth's structure et dynamics at lithospheric scales*



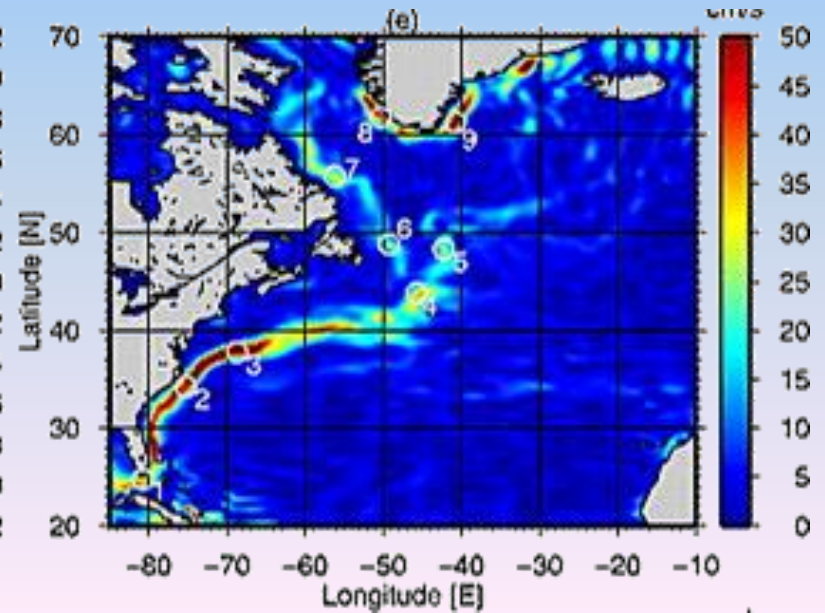
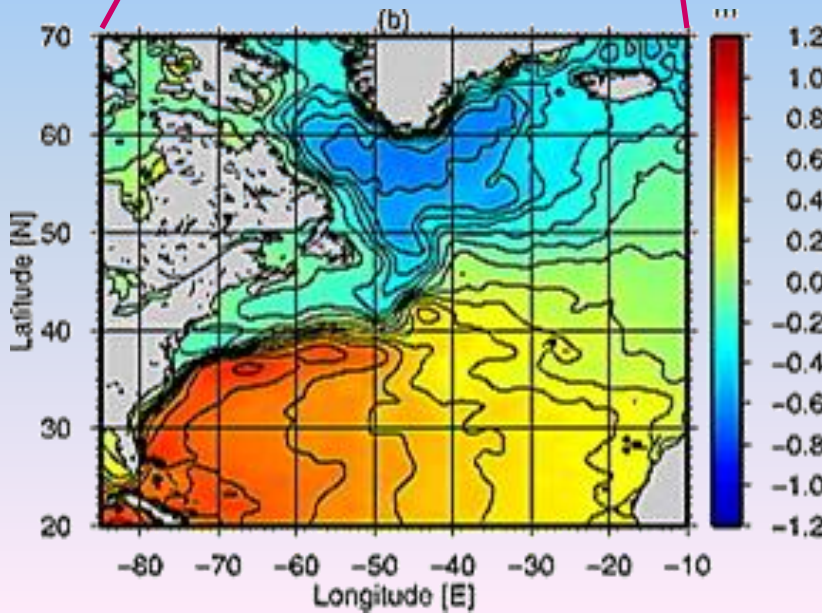
Example:  
dynamic  
topography of  
the oceans



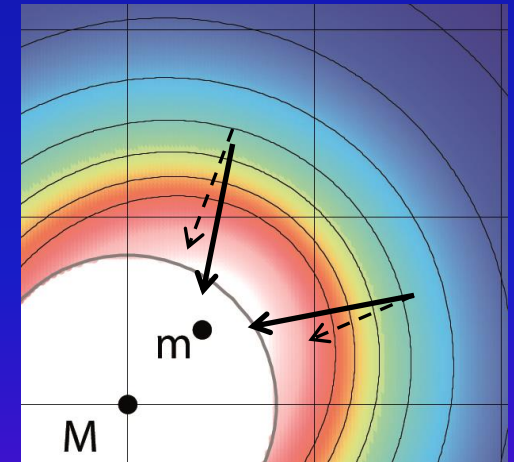
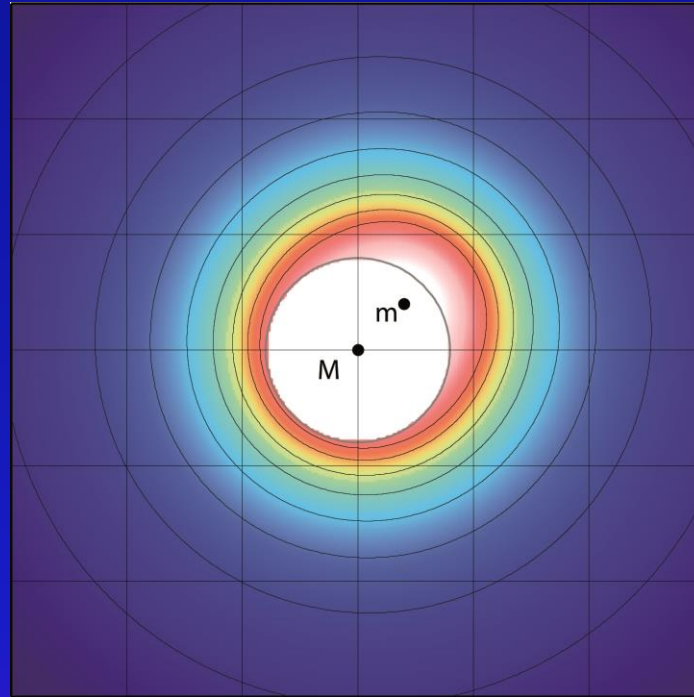
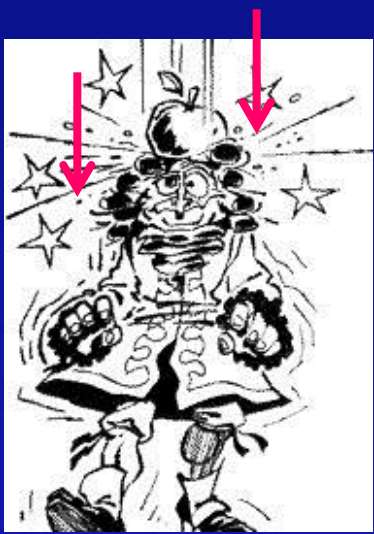
Arnault, 2004

Bingham *et al.* (2011) : dyn. topo. (m)

Geostrophic currents (cm/s)

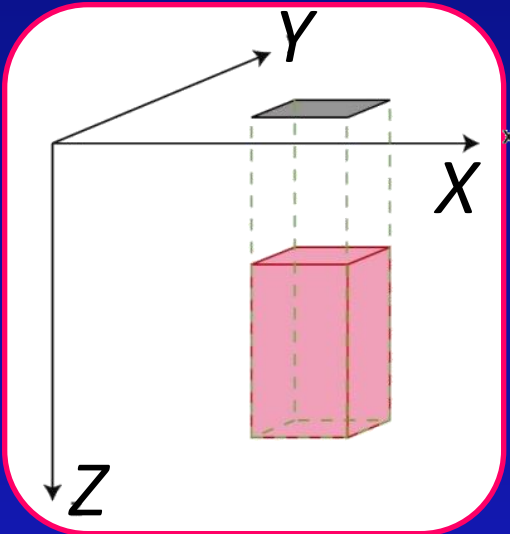


# Gravity is a vector

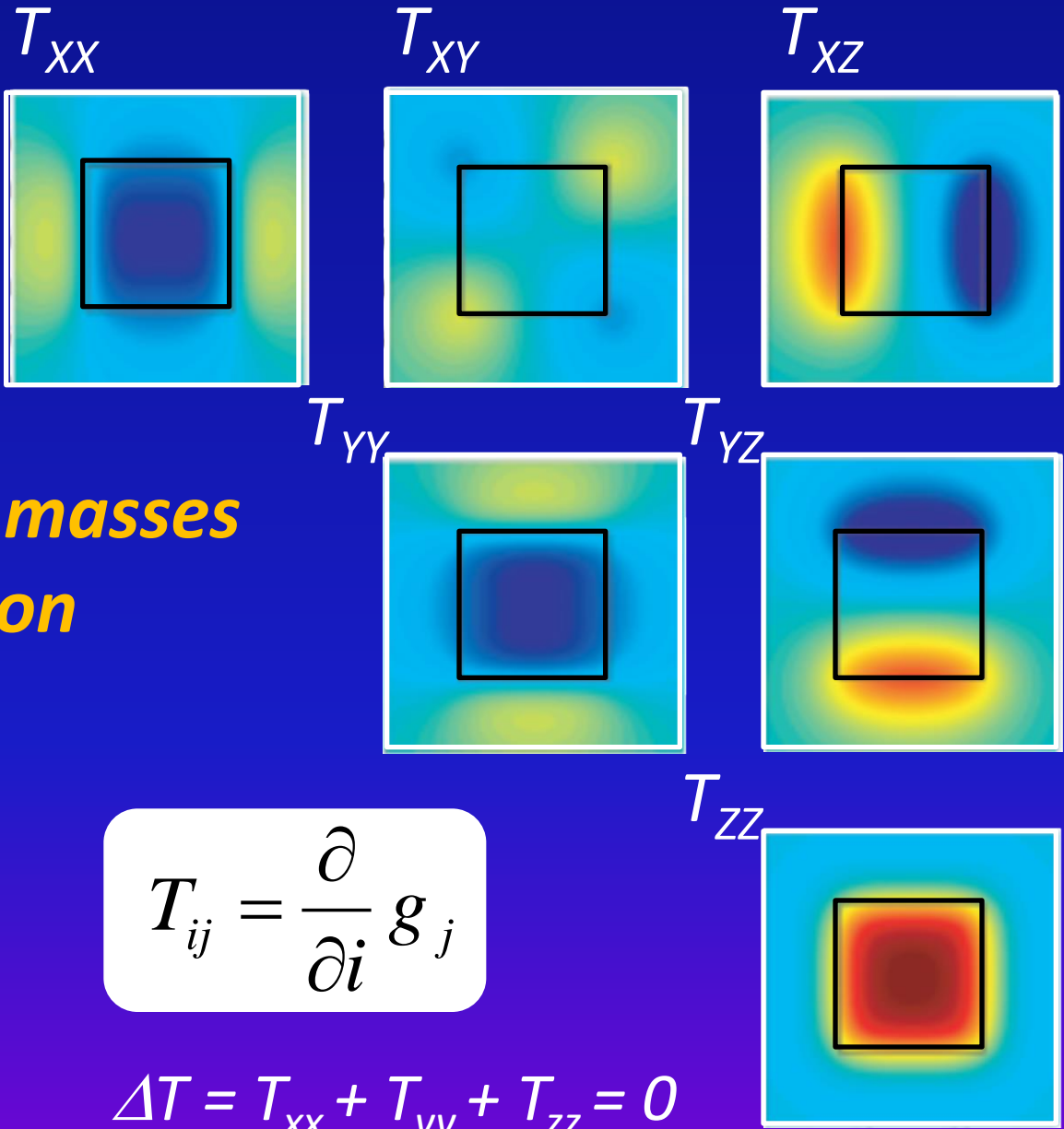


*Mass excess: locally, the gravitational attraction increases and its direction deviates towards the mass anomaly*



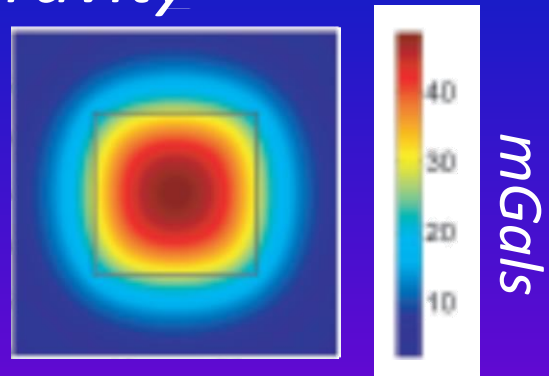


*Gradients tensor*



*Geometry of masses*  
*High resolution*

*Gravity*

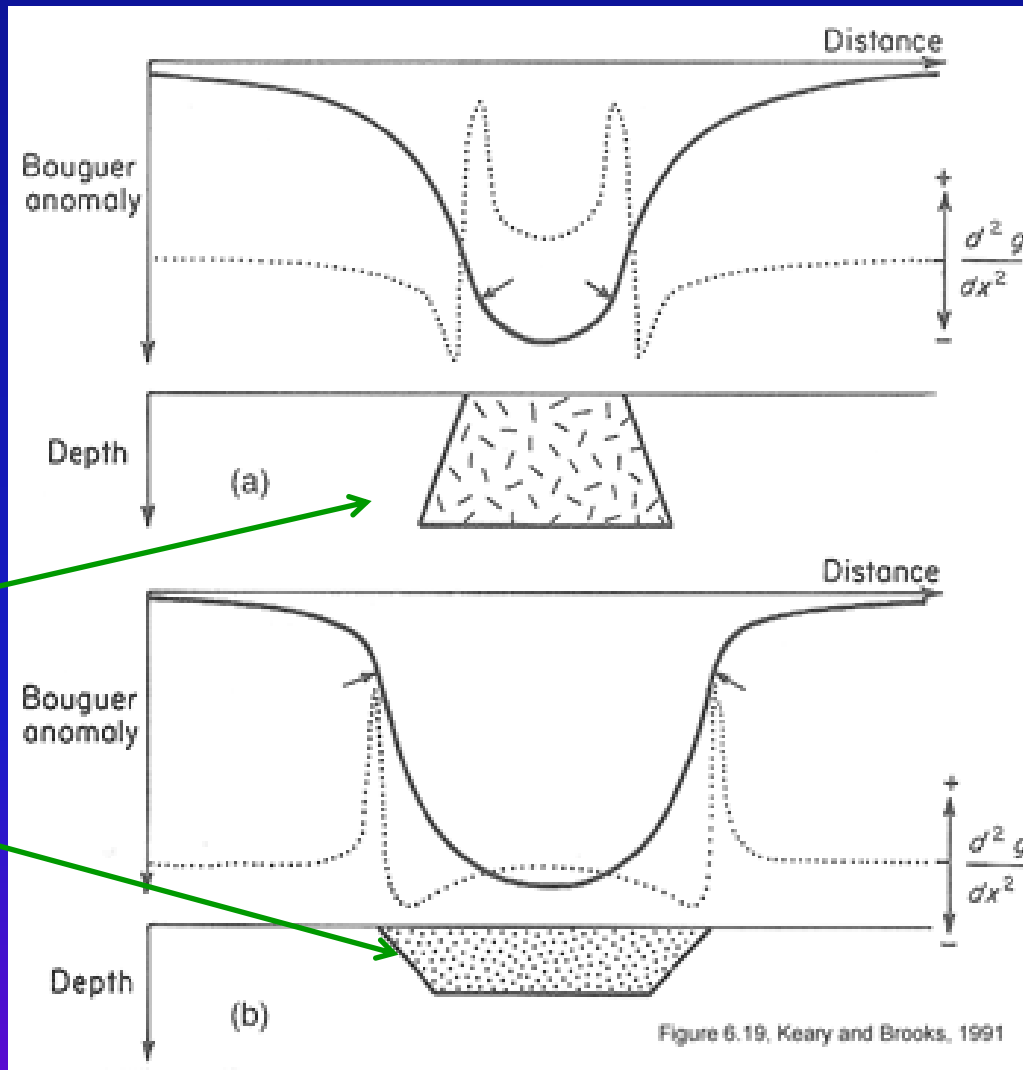


$$T_{ij} = \frac{\partial}{\partial i} g_j$$

$$\Delta T = T_{xx} + T_{yy} + T_{zz} = 0$$

*Images: Pajot (2008)*

# *Sensitivity of the horizontal gradients to the source geometry: another example*



mass  
anomaly

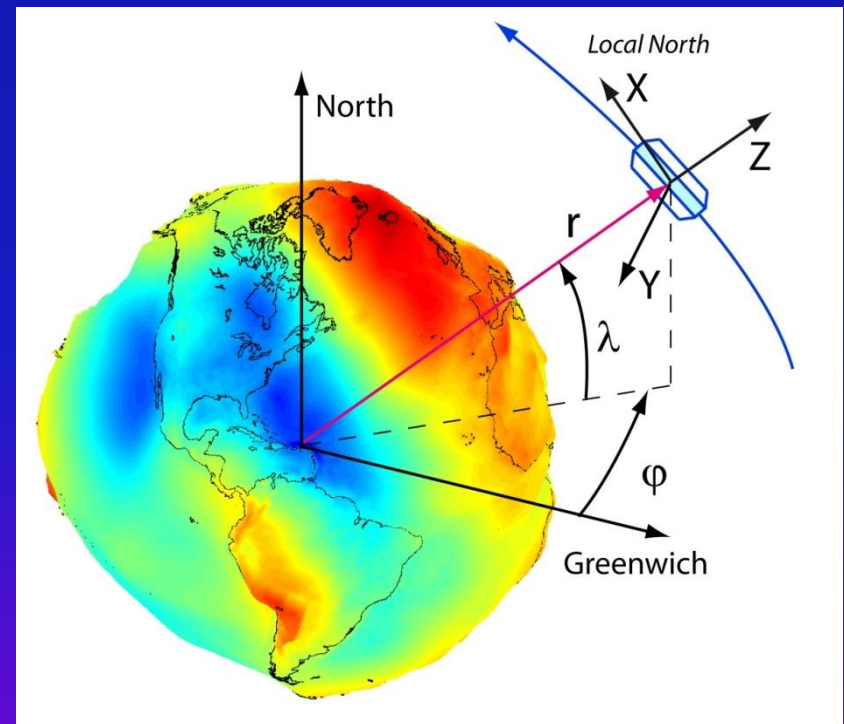
→ *used in exploration geophysics (local studies)*

# Earth's gravity gradients from satellites

Gradiometer data: scales  $< \sim 1000$  km

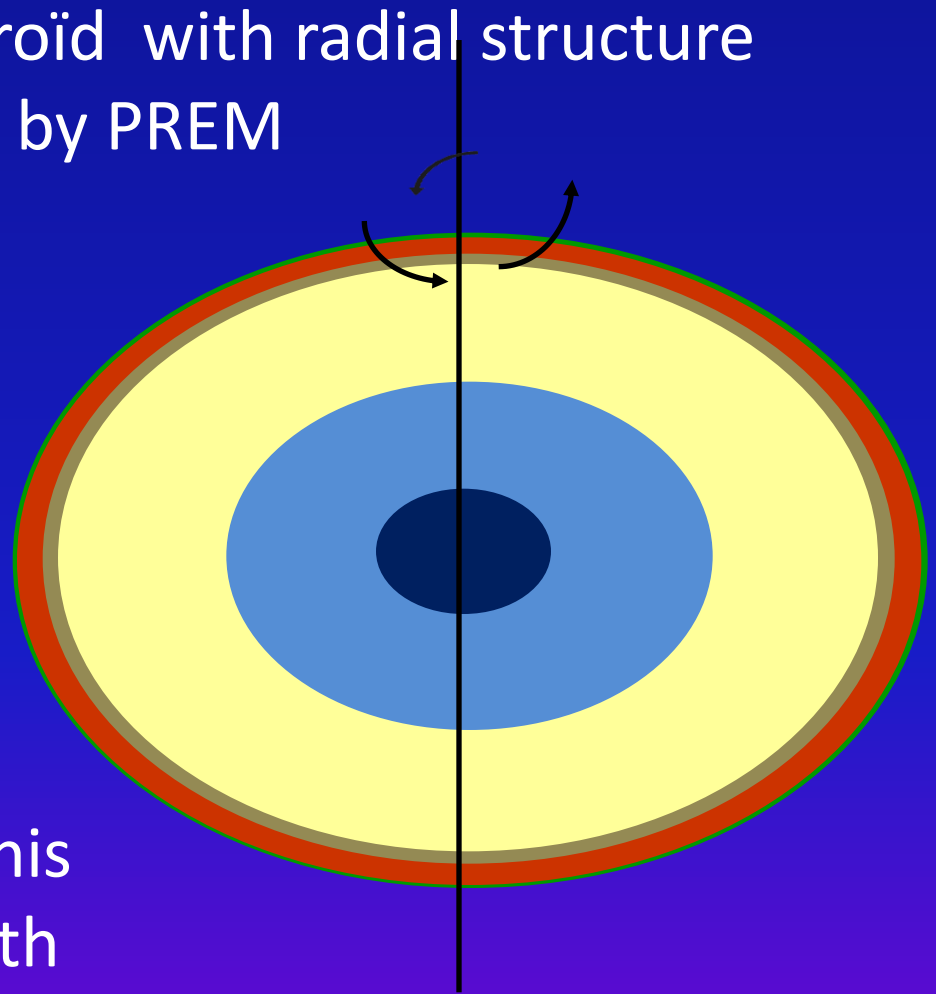
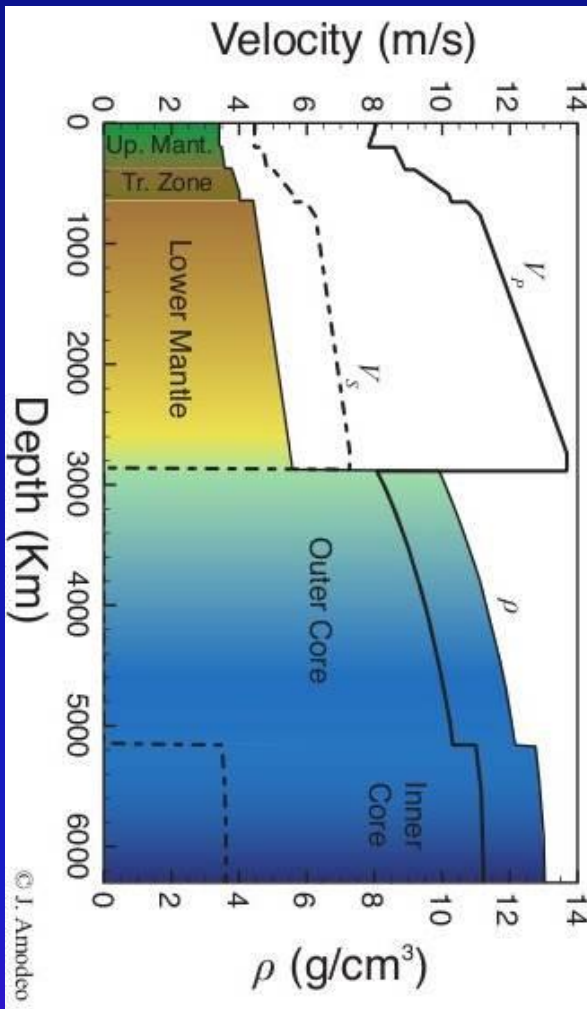
Period: Nov. 2009 - March 2011

Gradients expressed in the local  
North-oriented frame by the  
GOCE High Level Processing  
Facility



# Reference Earth's model

Spheroid with radial structure given by PREM

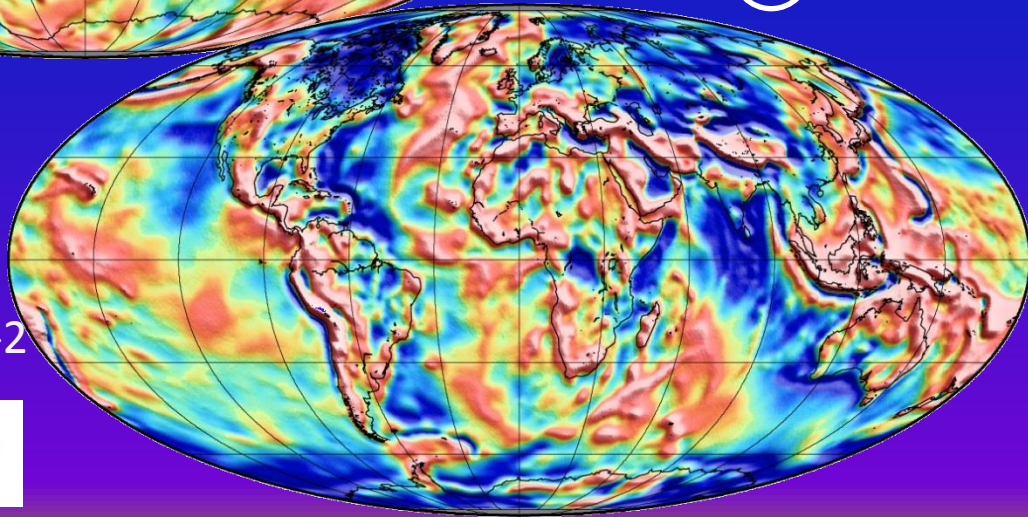
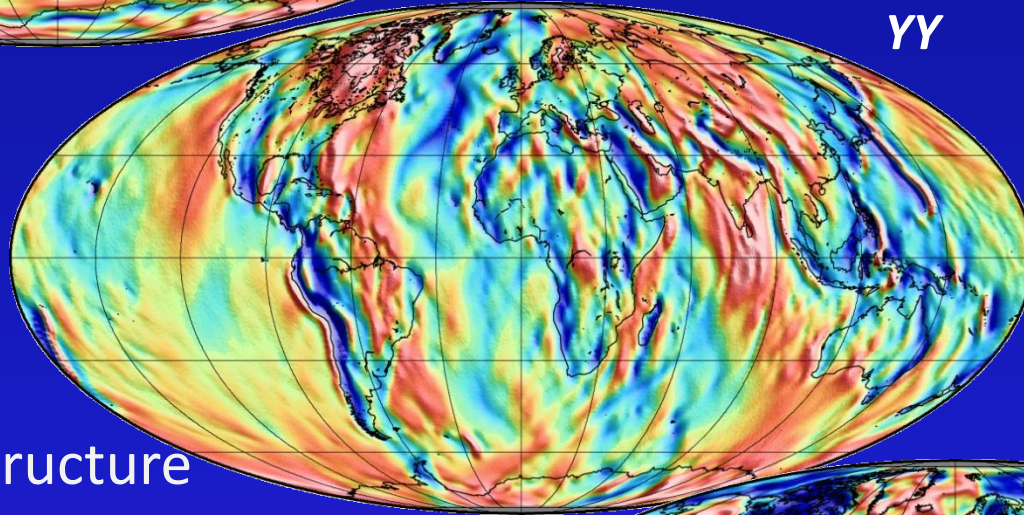
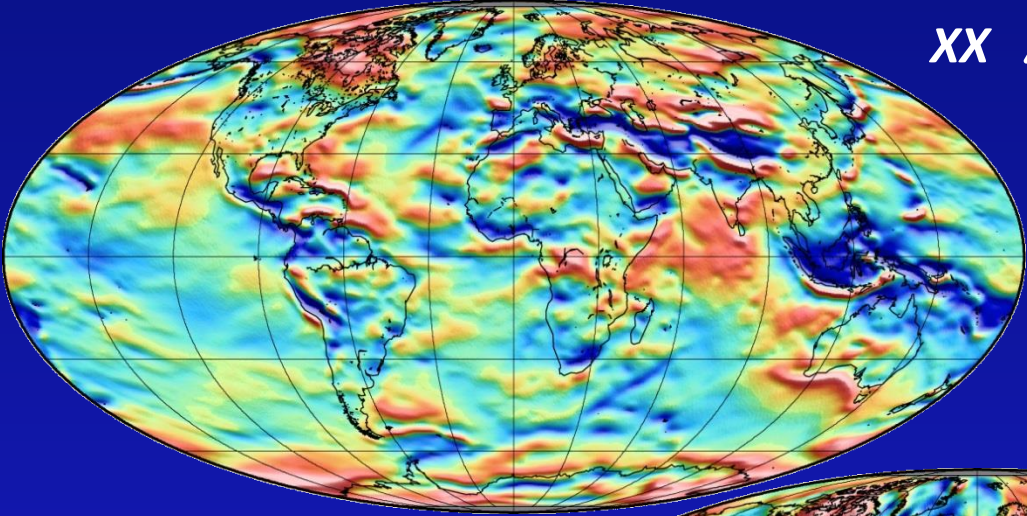


Hydrostatic equilibrium of this rotating, self-gravitating Earth

Chambat *et al.* (2010)



# Gradient anomalies at GOCE altitude



## Reference model:

- PREM radial structure
- Hydrostatic self-gravitating equilibrium of a rotating spheroid

$$1 \text{ Eötvös} = 10^{-9} \text{ s}^{-2}$$

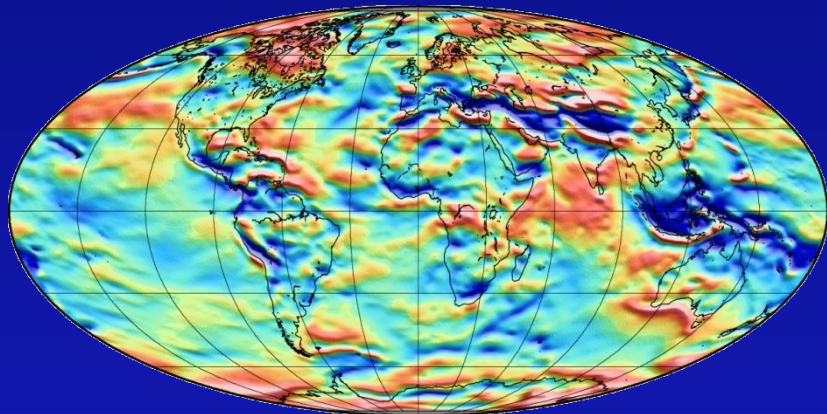




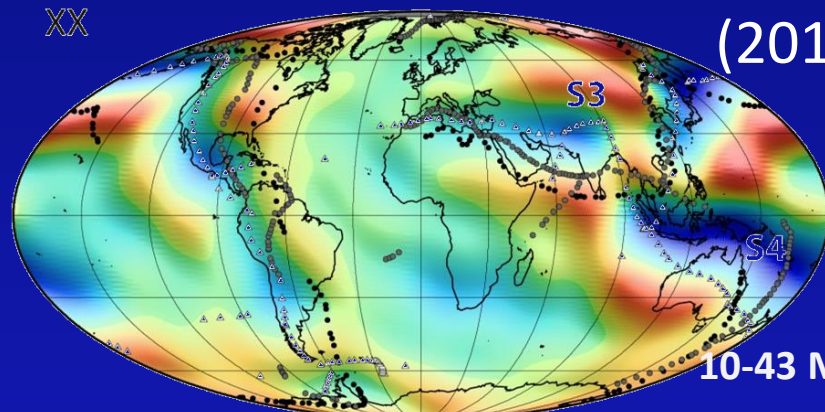
# Observed

# Modelled

using Rouby et al.  
(2010)



XX

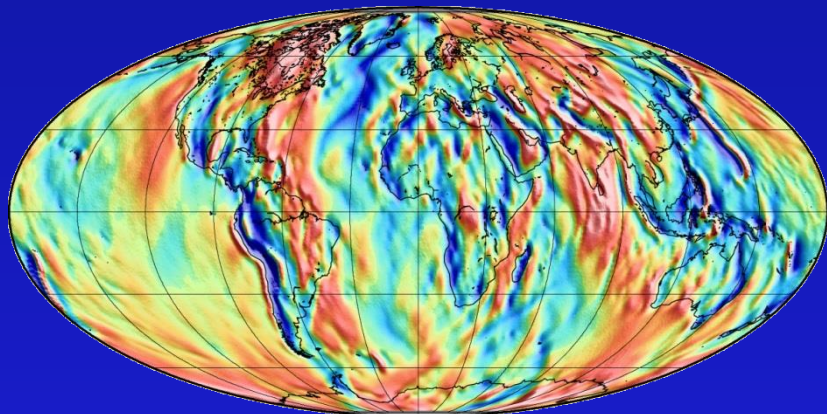


XX

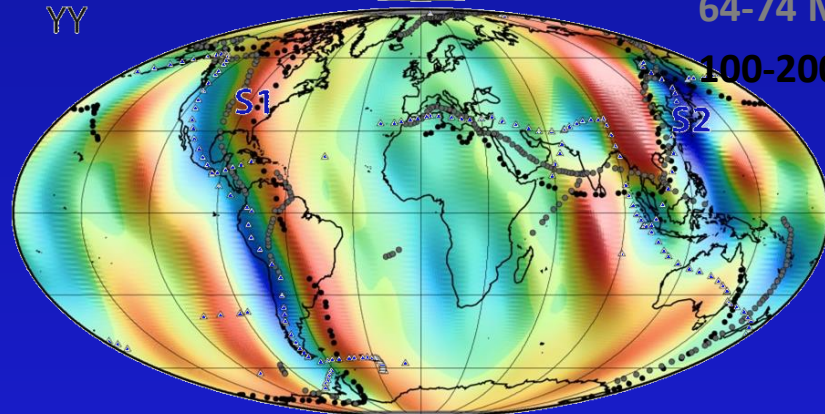
10-43 My

64-74 My

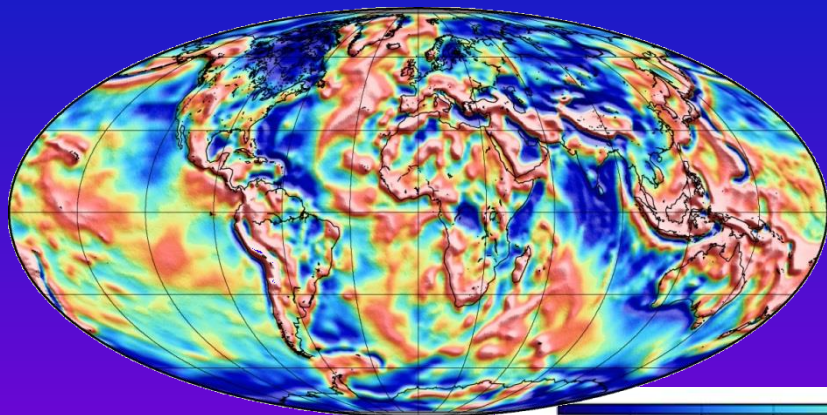
100-200 My



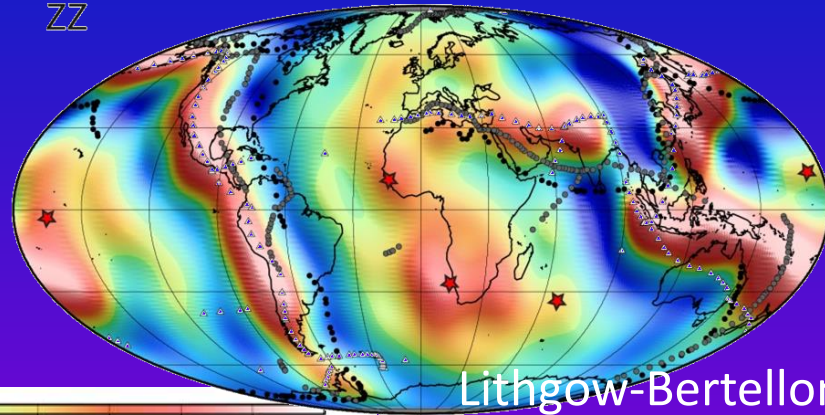
YY



YY

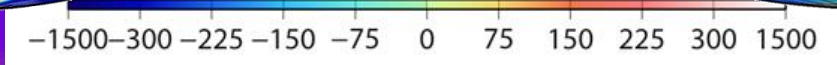


ZZ



ZZ

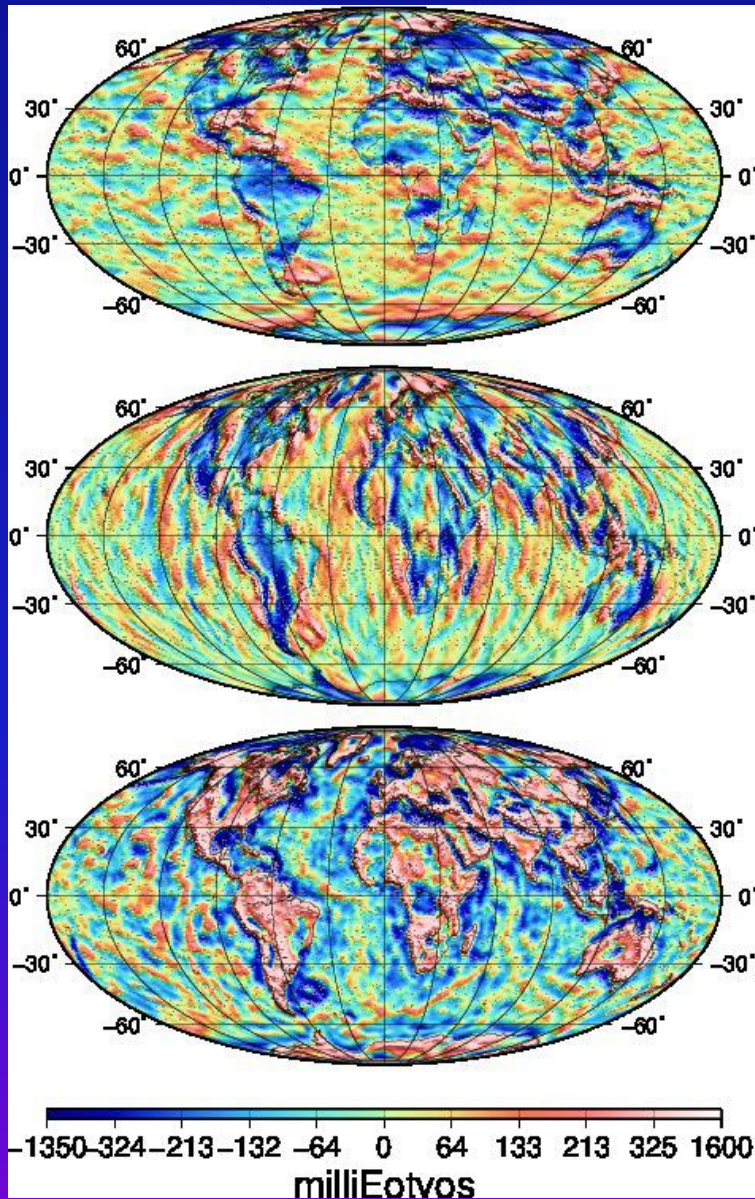
mEötvös



Lithgow-Bertelloni &  
Richards (1998)



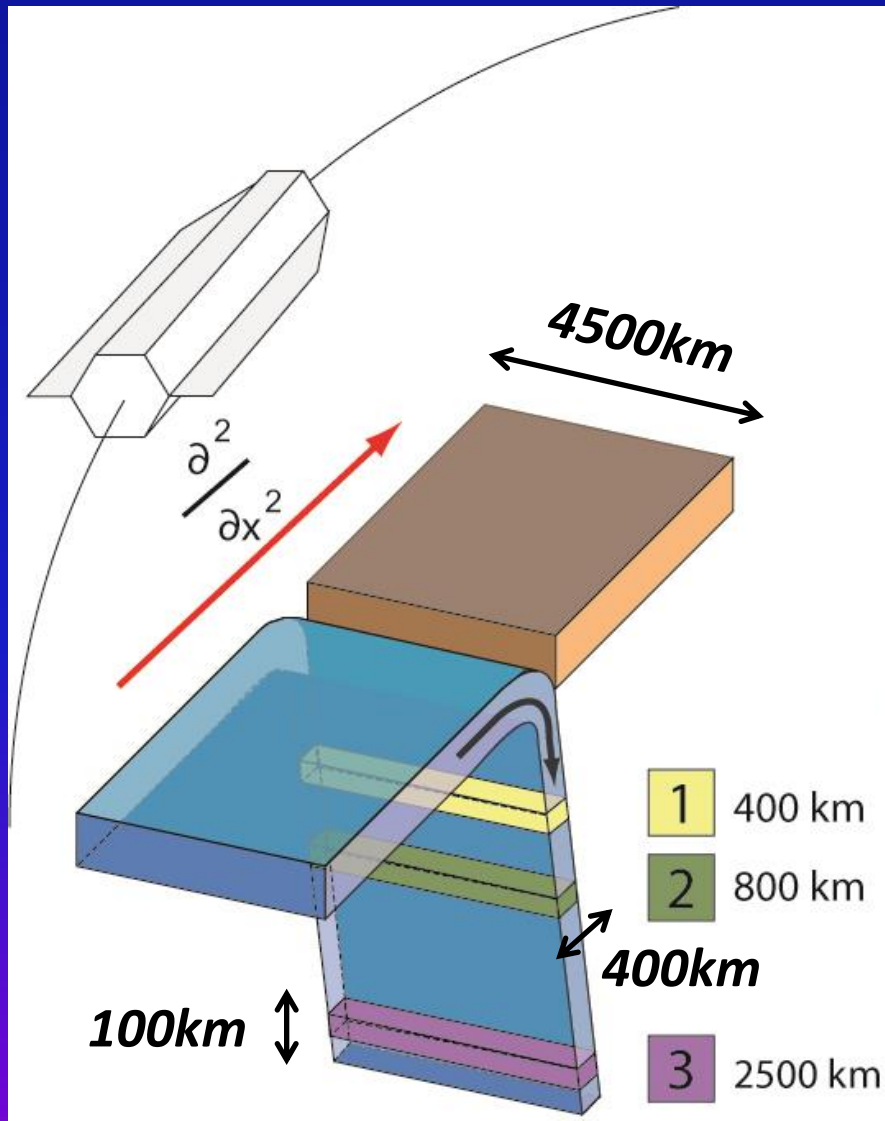
# Directionality helps separating sources



Gravity gradients associated to a simple crustal model at isostatic equilibrium

No such large-scale signal on  $XX$  and  $YY$

# What layers are probed and how?

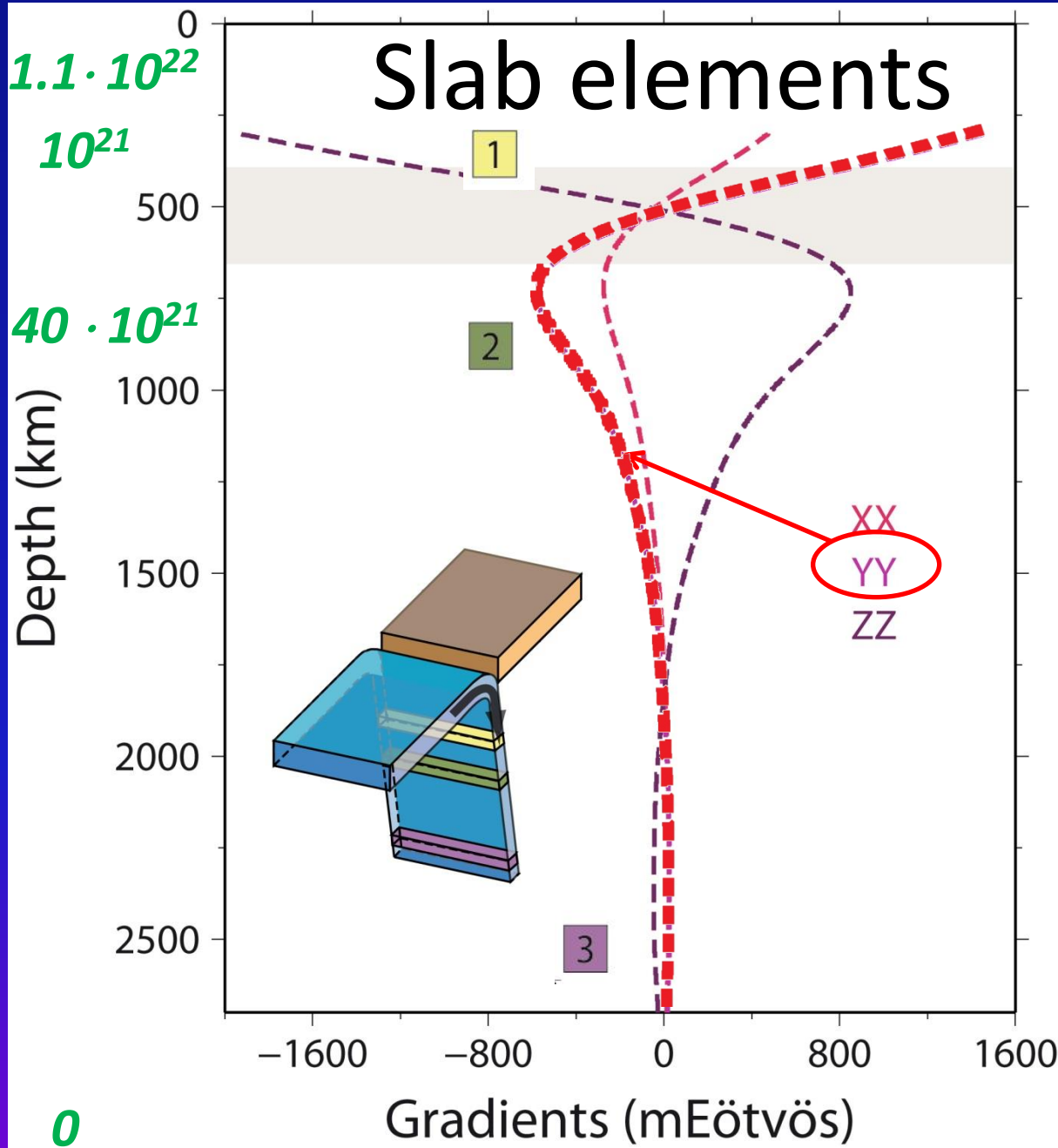


→ *Sensitivity analysis,  
example of slab elements*

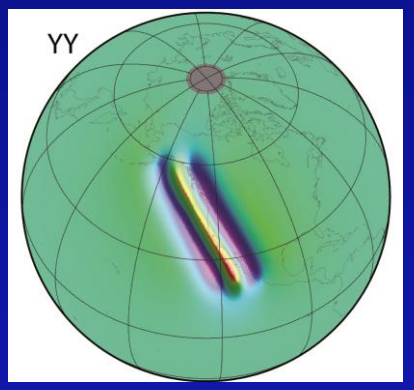
*Density contrast:  
+80 kg.m<sup>-3</sup>*



viscosity (Pa.s)

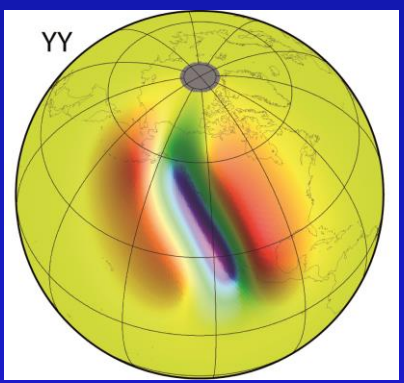


1

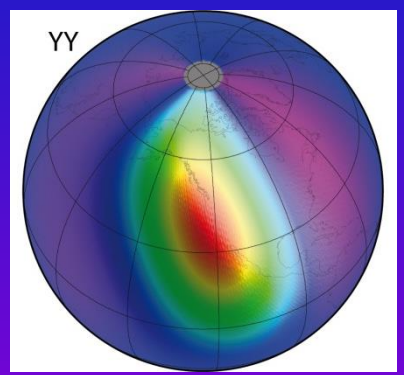


*oscillations at edges*

2



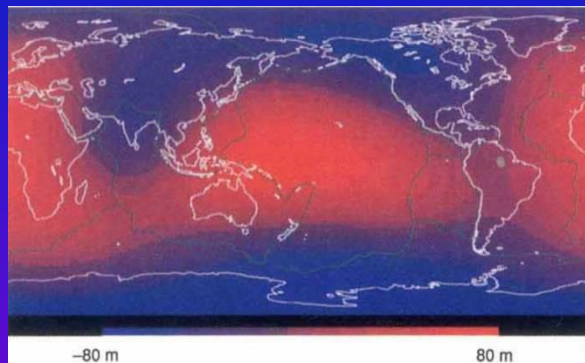
3



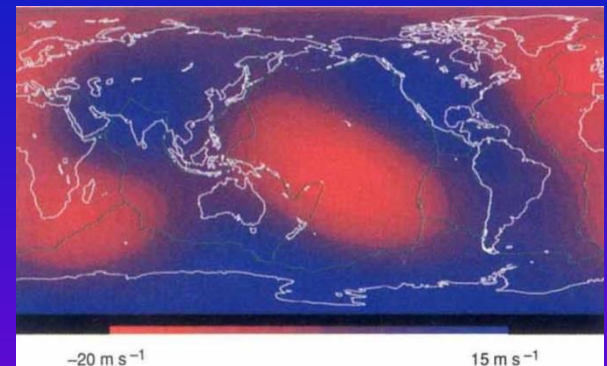
0

# Confrontation with seismic tomography

*First step: identification of common signals*

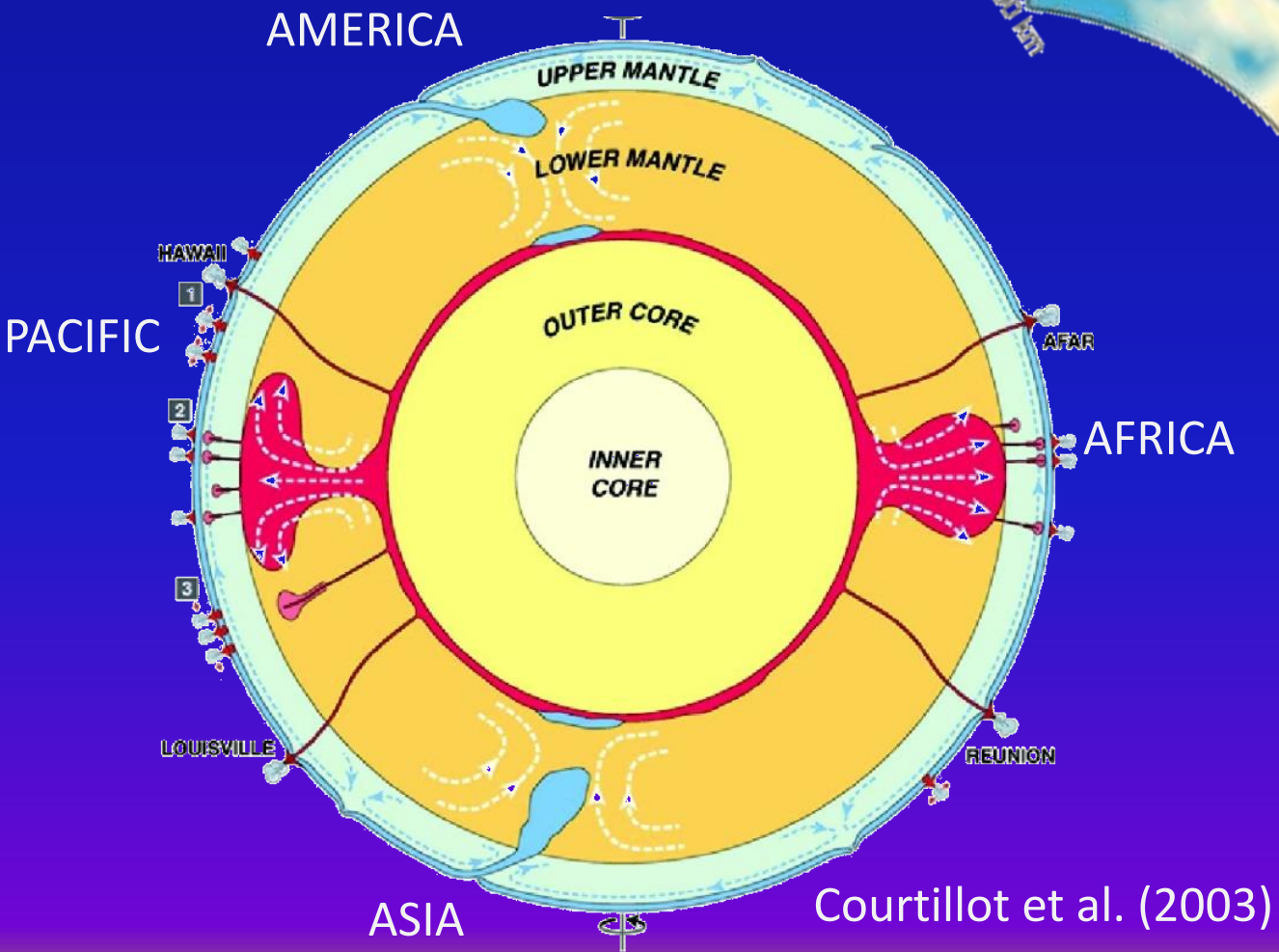


*Geoid*

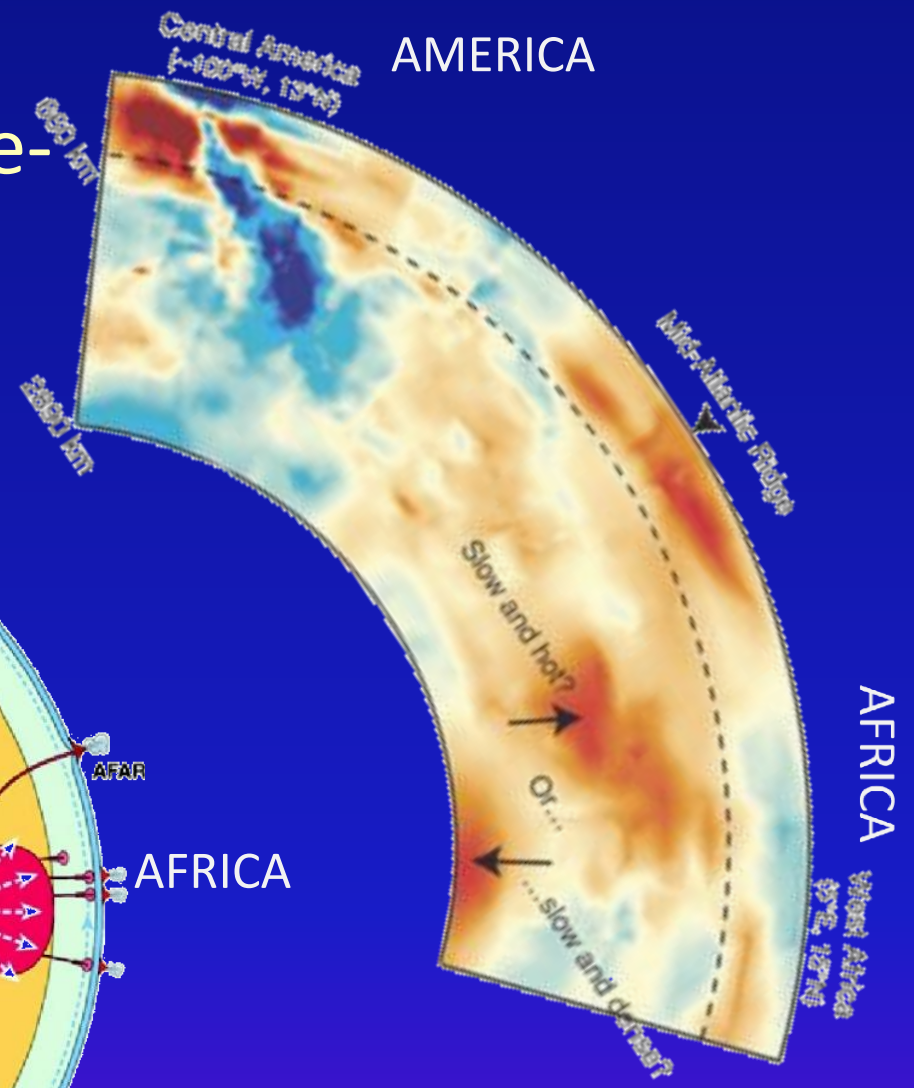


*dVs*

Interpreting seismic tomography in terms of time-varying dynamics requires independent data



Courtilot et al. (2003)



Van der Hilst (2004)



*Observed*

10-43 My

64-74 My

100-200 My

**YY gradients**

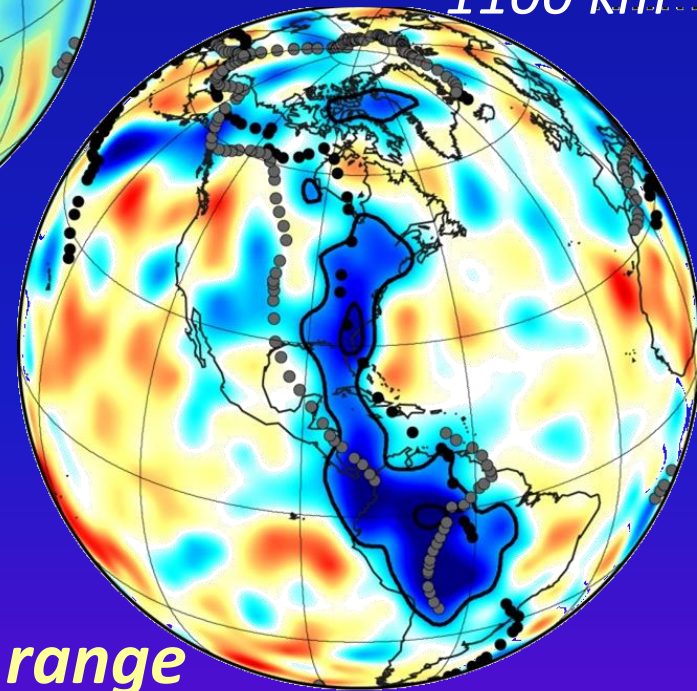
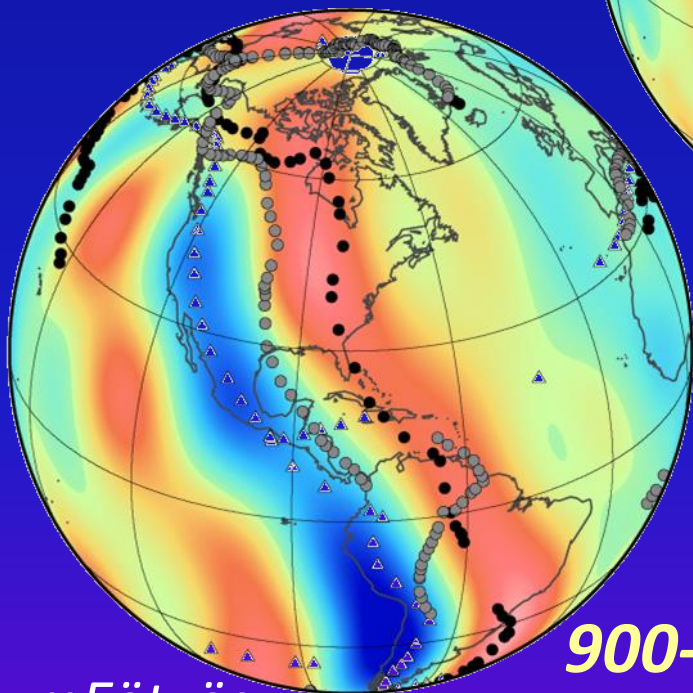
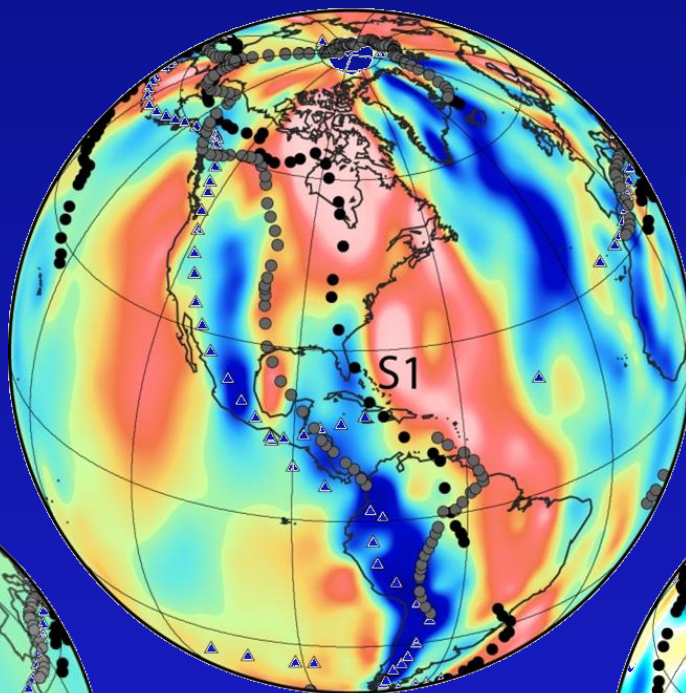
*Ritsema et al. (2011)*

**S4ORTS**

1100 km

*Modelled*

*Rouby et al. (2010)*

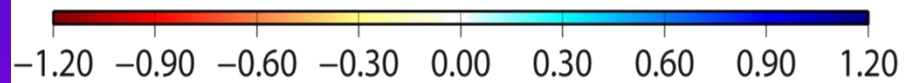
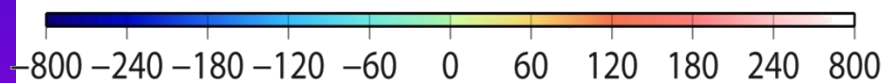


**Farallon slab**

**900-1600 km depth range**

*mEötvös*

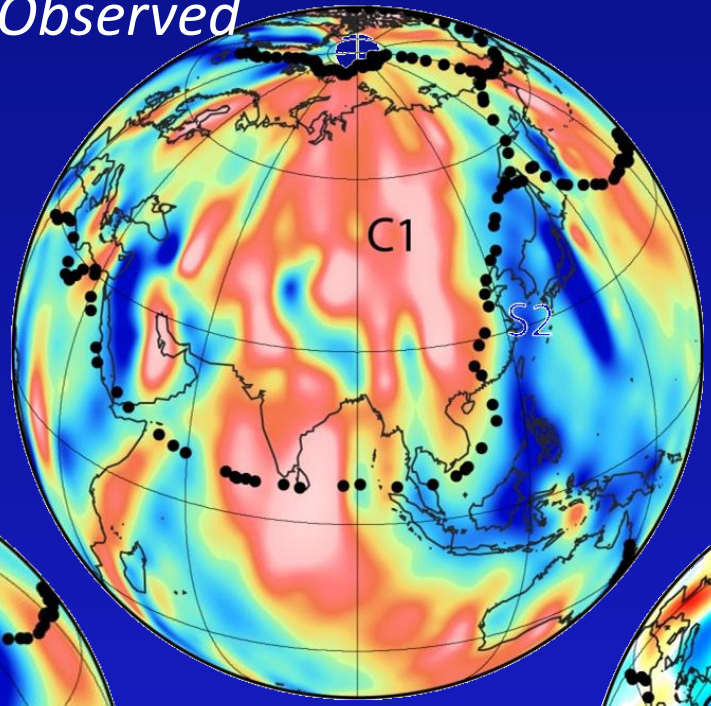
*dVs/Vs (%)*





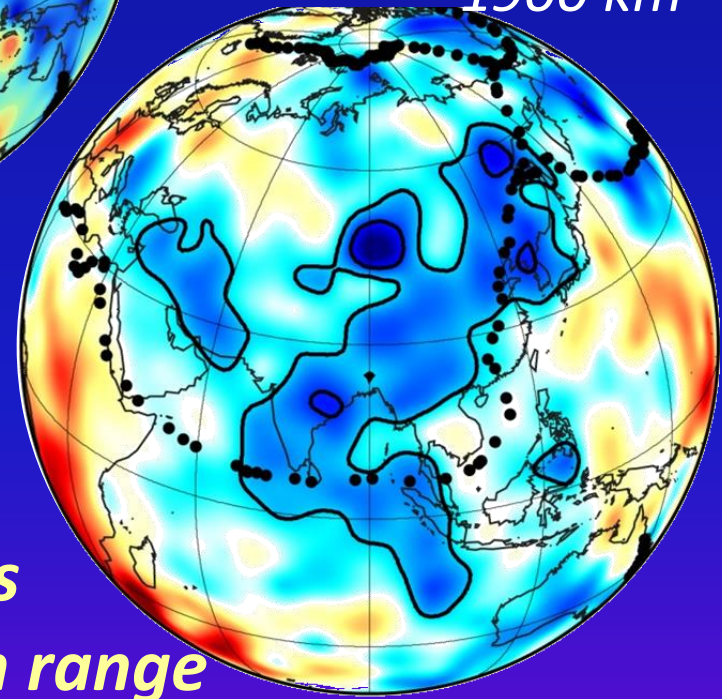
10-43 My  
64-74 My  
100-200 My

*Observed*



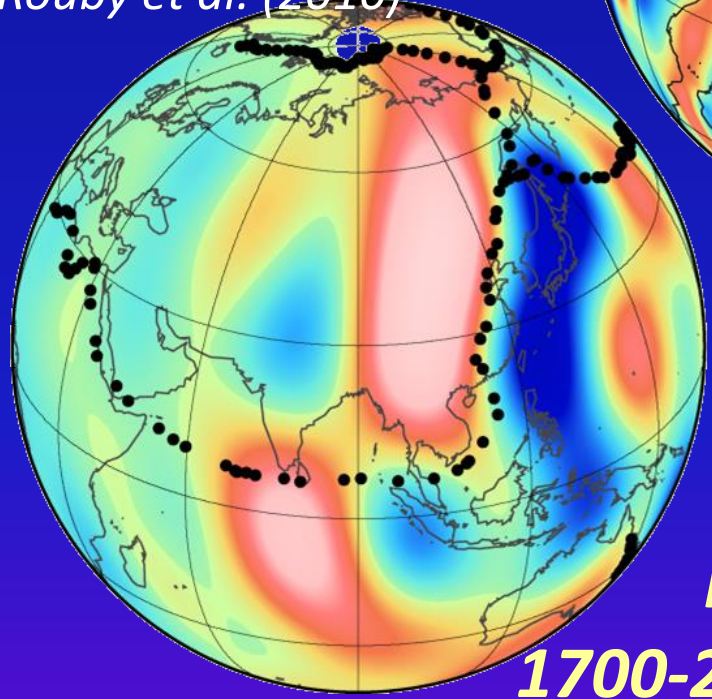
*Ritsema et al. (2011)*

**S4ORTS**  
1900 km



**YY gradients**

*Modelled*  
*Rouby et al. (2010)*

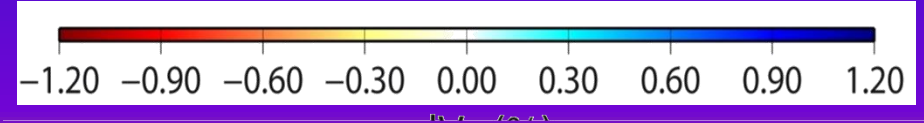


**Mesozoic slabs**

**1700-2600 km depth range**

*mEötvös*

*dVs/Vs (%)*



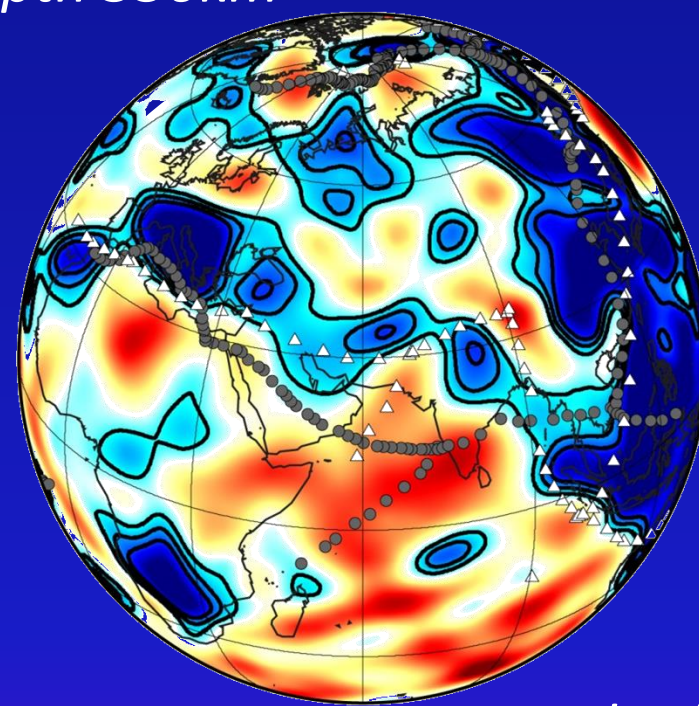
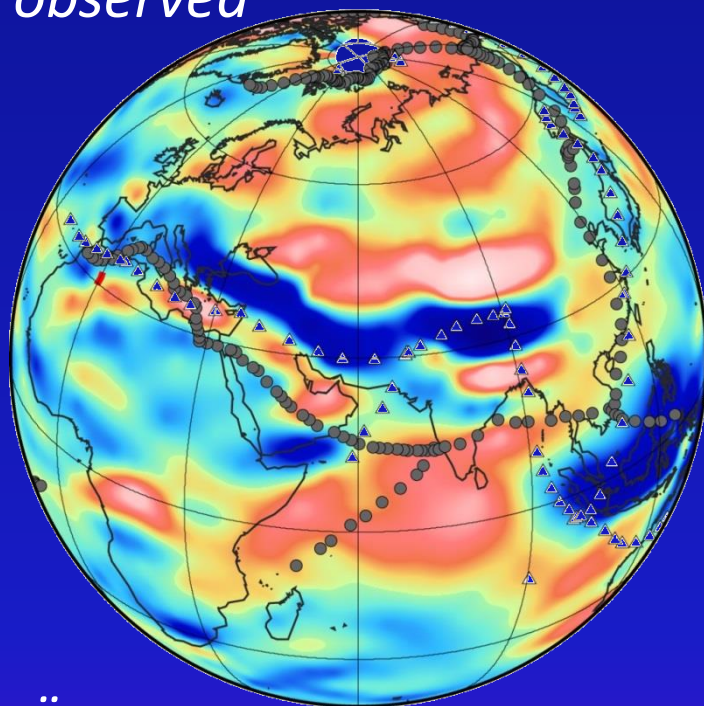
10-43 My

64-74 My

100-200 My

*XX gradients  
observed*

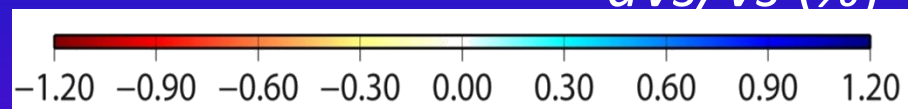
*DR2012    Debayle & Ricard (2012)  
depth 550km*



*mEötvös*



*dVs/Vs (%)*

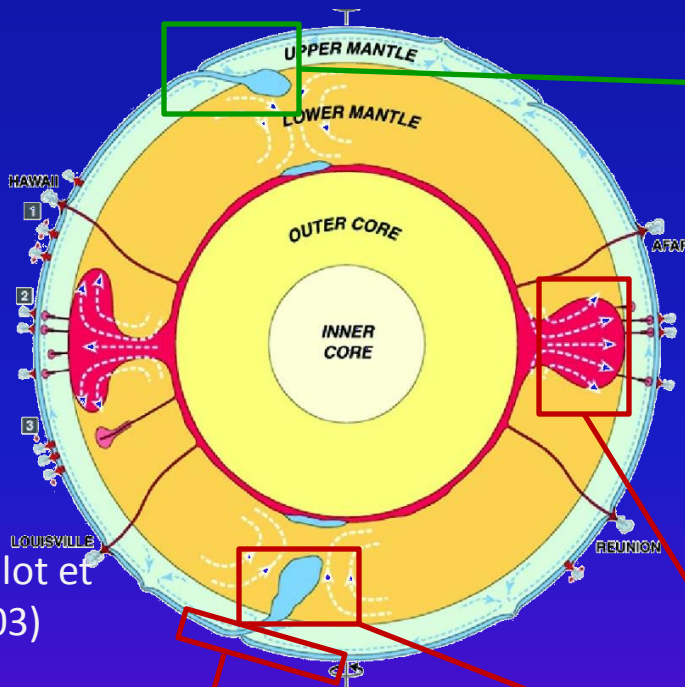


***E-W structure along the former Tethys - upper mantle?***

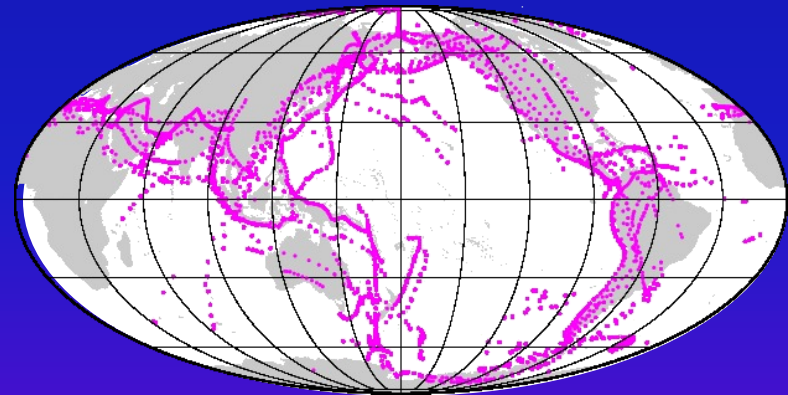


# Why do we detect so clearly the lower mantle contribution?

## *Strong sensitivity in the upper part of the lower mantle*



*Stability of almost North-South subductions around the Pacific over 250 M yr → the downwellings directionality coincides with that of the gradients*



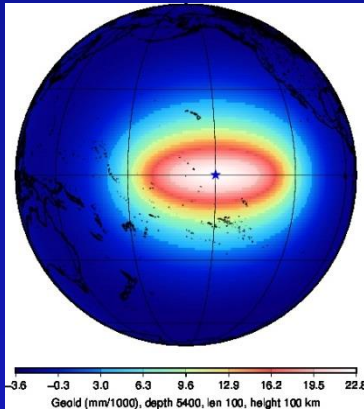
*Lithosphere signal reduced: strong sensitivity to isostasy*

*A lot of mass, not too much attenuation at satellite altitude*

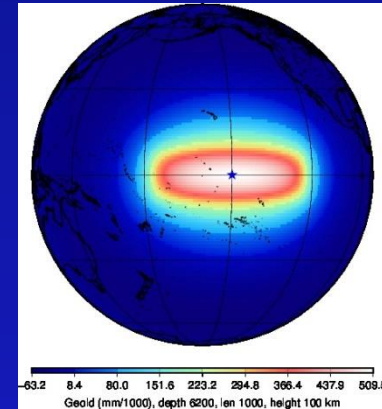
*Thin and deep*

or

*Wide and shallow?*

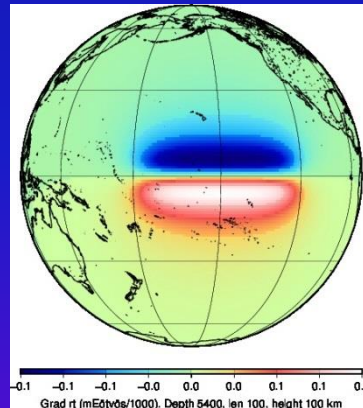
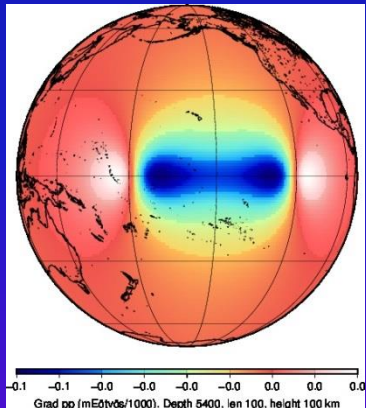


*Geoid*



$T_{PP}$

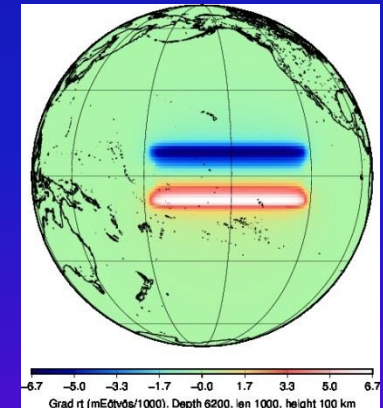
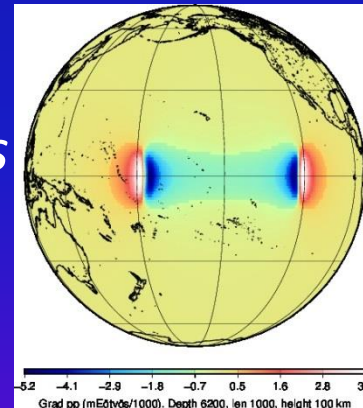
$T_{RT}$



*Gradients*

$T_{PP}$

$T_{RT}$



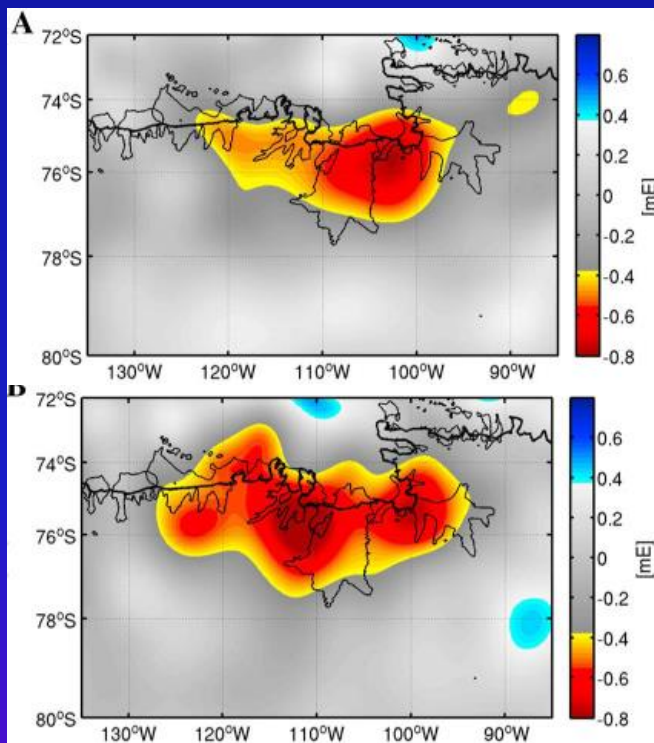
*Less ambiguity than classical gravity*

*→ more efficient combination to seismology*



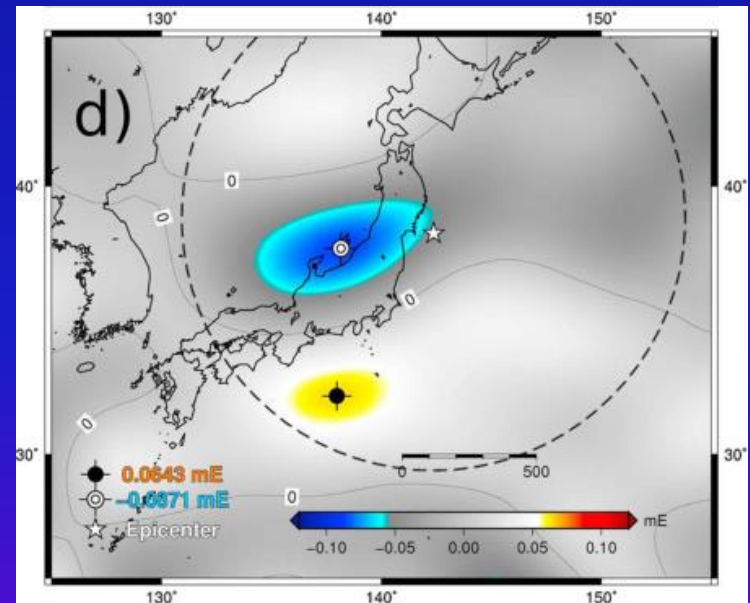
# Time variations of GOCE gravity gradients?

- Even if GOCE was not intended for it, detection of local slow/long term gradients variations  $\leq$  mEötvös



03-06/2010

07-09/2011



*Evolution of ice depletion signal in Amundsen Sea Sector, Antarctica (Bouman et al., GRL, 2014)*

*Tohoku 2011 earthquake signal (Fuchs et al., JGR, 2013)*

# Conclusions and outlooks

- From GOCE: interest of a vectorial view of gravity to image masses geometries, not only at small scales (unexpectedly). Geometric consistency with seismology, which does not give the mass, makes the combination possible.
- Large scales rely on the orbit or on GRACE data → no measurement of large-scale gravity gradients.
- Another unexpected result comes from time-varying gravity gradients. Identification of slow signals at high spatial resolution (ice mass loss, giant earthquake).

# Conclusions and outlooks

- Strength of gravity gradients to separate superimposed sources based on shapes and directions
- This is true not only for the quasi-static components of the gravity field (geological time scales), but also for its faster time variations...
  - *Separation between solid Earth and climate signals*
- Gravity gradients from Microscope could bring new insights on Earth system dynamics